

Surface active substances in the upper water column during a Southern Ocean Iron Fertilization Experiment (EIFEX)

P. L. Croot,¹ U. Passow,² P. Assmy,² S. Jansen,² and V. H. Strass²

Received 5 September 2006; revised 10 January 2007; accepted 18 January 2007; published 15 February 2007.

[1] Surface active substances (SAS) in the water column were measured by voltammetry using the electrochemical probe *o*-nitrophenol (ONP) during EIFEX, a mesoscale open ocean iron enrichment experiment in the Southern Ocean. SAS levels were low throughout the experiment (<0.005 – 0.03 mg L⁻¹ Triton X-100 equivalents). Initially SAS was extremely low in the photic zone, but as the phytoplankton bloom developed concentrations markedly increased throughout the upper 100 m (~0.02 mg L⁻¹ Triton X-100 equivalents). Highest concentrations of SAS (>0.02 mg L⁻¹ Triton X-100 equivalents) were found at the end of the bloom particularly at density discontinuities where organic material may accumulate. Exudates from diatoms appeared to be the major source of SAS during EIFEX, either from direct extracellular release or in the action of being grazed upon by zooplankton. **Citation:** Croot, P. L., U. Passow, P. Assmy, S. Jansen, and V. H. Strass (2007), Surface active substances in the upper water column during a Southern Ocean Iron Fertilization Experiment (EIFEX), *Geophys. Res. Lett.*, 34, L03612, doi:10.1029/2006GL028080.

1. Introduction

[2] Amongst the wide variety of dissolved organic molecules present in seawater are an important subset with surface active properties [Hunter and Liss, 1981]. These molecules typically possess structural groups with affinity for water while also containing other structural groups that are hydrophobic in nature; as such these molecules tend to accumulate by adsorption processes at water phase boundaries. In this manner SAS influences numerous geochemical and physical processes in the ocean. The presence of SAS in the sea surface micro-layer can significantly decrease air-sea exchange rates of dissolved gases [Frew *et al.*, 1990; Tsai and Liu, 2003]. Surfactants can also alter the equilibrium between the dissolved and particulate phases for many trace metals [Shine and Wallace, 1996] and enhance the aggregation of small colloids and particles, leading to increased sinking fluxes [Mopper *et al.*, 1995; Zhou *et al.*, 1998]. The source of these surfactants is presumed to be primarily from phytoplankton exudates and their degradation products either by direct production [Zutic *et al.*, 1981] or from zooplankton grazing [Kujawinski *et al.*, 2002] with large concentrations of surfactants often being observed during phytoplankton

blooms in coastal waters [Gašparović and Čosović, 2001; Gašparović *et al.*, 1998a]. The composition of the surface-active material is strongly influenced by the dominant phytoplankton [Gašparović *et al.*, 1998b] and is mostly comprised of polysaccharide compounds [Passow *et al.*, 1994] but also includes biologically derived lipids and proteins.

[3] In the present work samples from the upper water column (0–300 m) were measured for SAS during the course of the European Iron Fertilization Experiment (EIFEX: R/V *Polarstern* cruise ANT XXI/3) in the Atlantic sector of the Southern Ocean from 9 Feb to 21 March 2004 [Smetacek *et al.*, 2005]. This work was performed to complement other data sets collected during EIFEX (papers in preparation) on: (1) transparent exopolymer particles (TEP), in which SAS may be precursor compounds [Passow, 2000; Zhou *et al.*, 1998]; and (2) trace metal data, where SAS may affect the removal rates and speciation of certain metals in the water column [Wallace, 1982] and in the sea surface microlayer [Hardy *et al.*, 1985; Hunter, 1980]. This work represents the first measurements of SAS during an open ocean iron induced phytoplankton bloom.

2. EIFEX Setting

[4] The EIFEX study was performed in a mesoscale cyclonic eddy, embedded in a meander of the Antarctic Polar Front [Strass *et al.*, 2005]. The eddy was approximately 90 × 120 km in scale centered on 49.4° S, 2.25° E. The initial fertilization was performed on 12–13 February 2004 by the addition of 7000 kg of iron sulfate (FeSO₄ · 7H₂O) as an acidic slurry into the mixed layer. The fertilized area was a spiral around the eddy centre, corrected for lagrangian drift, as determined by a GPS-buoy released immediately prior to the release of the Fe. The initial fertilization area was approximately 150 km². After two weeks with dissolved Fe concentrations approaching pre-infusion levels a second infusion of 7000 kg of iron sulfate took place within the previously enriched patch as defined by the photosynthetic activity Fv/Fm [Röttgers *et al.*, 2005] as measured by Fast-Repetition-Rate-Fluorescence (FastTracka, Chelsea, UK). All sampling sites reported in this manuscript were located in the eddy; stations within the fertilized waters are hereafter described as in-patch and those from outside as out-patch. Identification of stations as being in or out of the fertilized patch was assessed based on several key criteria: iron concentration, Fv/Fm, pCO₂ and chlorophyll [Smetacek *et al.*, 2005].

3. Methodology

[5] Electrochemical methods utilizing *o*-nitrophenol (ONP) as a probe for the rough characterization of organic

¹Forschungsbereich Marine Biogeochemie, Leibniz-Institut für Meeresswissenschaften at University of Kiel, Kiel, Germany.

²Alfred Wegener Institute for Polar and Marine Research, Bremerhaven, Germany.

Table 1. Station Information for EIFEX Surfactant Sampling

Station	Date ^a	Time, UTC	Latitude	Longitude
425 - Fe Infusion	12/2/2004 (0)	13:15	49.4° S	2.25° E
427 IN/OUT ^b	13/2/2004 (1)	19:05	49° 12.01' S	02° 04.99' E
508	22/2/2004 (10)	07:53	49° 12.14' S	02° 00.44' E
512 - Fe Infusion	26/2/2004 (14)	12:31	49° 38.7' S	02° 30.1' E
514 OUT	29/2/2004 (17)	17:53	49° 15.42' S	02° 20.14' E
543	04/3/2004 (21)	06:54	49° 28.76' S	02° 27.32' E
570	14/3/2004 (31)	02:17	49° 25.74' S	02° 03.17' E
580	16/3/2004 (33)	02:28	49° 07.88' S	02° 15.21' E

^aThe number in parentheses indicates the number of days after the first infusion began, the numbers are rounded to the nearest integer value.

^bStation 427 was sampled immediately after the initial Fe fertilization and is thus used here as the initial time point (data not shown in Figure 1, SAS concentrations were at or below the detection limit: SAS < 0.005 mg L⁻¹ Triton X-100 eq).

matter in seawater have been developed in a series of papers by Gašparović, Čosović and Vojvodić [Gašparović and Čosović, 1994, 1995, 2001; Gašparović et al., 1997, 1998a, 1998b]. In the present work SAS was determined with Alternating Current -Voltammetry (in-phase mode) with a Metrohm VA 757 voltammeter using ONP as a probe compound [Gašparović et al., 1998a]. Standards, using Triton X-100 as the external standard, were run in samples of deep seawater (1500 m - low surfactants) which had previously been UV oxidized (Metrohm UV-Digestor 705) to remove organic material (see Text S1 in the auxiliary material).¹ To avoid contamination from surfactants present in the sea surface microlayer, samples were collected from GO-FLO (General Oceanics Model 1080) sampling bottles deployed on a trace metal clean hydrowire (Aramid). Filtered (acid-cleaned 0.2 μm Sartorius cartridge filter) samples were drawn into ethanol rinsed polyethylene vials and kept at 4° C (in situ temperature) in the dark until analysis. Aliquots (20 mL) of the seawater samples were spiked with ONP in solvent (HPLC grade - ethanol) cleaned Quartz or Teflon cell cups. The samples were purged with dry N₂ gas for 3 minutes and pre-concentrated on the Hg drop for deposition times of 0, 60, 180 and 300 s.

[6] Hydrographic data from each station was collected from separate conductivity-temperature-depth (CTD) casts using a Seabird SBE911plus (Seabird Electronics) with a 24-bottle water sampler (General Oceanics). The instruments were calibrated immediately before and after the cruise. For in situ calibration, temperatures were measured with reversing thermometers, and salinity samples were analyzed with a Guildline-Autosal-8400A salinometer onboard.

[7] The composition of the phytoplankton community during EIFEX was assessed by light microscopy. Water samples of 200 mL were obtained from the CTD rosette from 7 depths between 10 and 150 m for the quantitative assessment of the diatom assemblage. Samples were preserved with hexamine-buffered formaldehyde solution (2%) and stored at 4° C in the dark for subsequent counting in the home laboratory. A volume of 50 mL was settled in sedimentation chambers (Hydrobios, Kiel, Germany) for 48 hours. Cells were identified and enumerated using inverted light and epifluorescence microscopes (Axiovert 135 and Axiovert 200, Zeiss, Oberkochen, Germany) according to Utermöhl [1958]. Bio-volume was calculated from equivalent geometrical shapes [Hillebrand et al., 1999] and converted to cellular carbon content through recom-

mended carbon conversion equations [Menden-Deuer and Lessard, 2000]. Biomass of all species described here are trapezoidal depth-integrated values for the upper 0-100 m of the water column.

4. Results and Discussion

4.1. Water Column Surfactants During EIFEX

[8] Samples were collected and analyzed onboard during EIFEX from 6 stations during the course of the phytoplankton bloom (Table 1 and Figure 1). The concentration of SAS was extremely low initially at both the in and the out stations throughout the water column (<0.005 mg L⁻¹ Triton X-100 equivalents). SAS was lower than published data from the Adriatic [Gašparović and Čosović, 2001] and sub-Arctic Fjords [Gašparović et al., 2005], as might be expected for an open ocean, deep mixed layer, low chlorophyll location. SAS levels increased sharply at depth around day 10, prior to the 2nd infusion (day 14) but were still low outside the patch (day 17). As the experiment progressed there was an increase in SAS (~0.02 mg L⁻¹ Triton X-100 eq) throughout the upper water column (day 21). For the last week of observations of the EIFEX patch, enhanced levels (~0.03 mg L⁻¹ Triton X-100 eq) of surfactant activity were found in the water column immediately below the shallowing active mixed layer (AML) (as defined in Brainerd and Gregg [1995]), similarly enhanced levels were also found at depths below the mixed layer (Figure 1).

4.2. Development of the EIFEX Bloom

[9] During the course of EIFEX chlorophyll *a* concentrations increased roughly 6 fold up to 26 days after the initial infusion and then began to decrease [Hoffmann et al., 2006]. The Fe addition saw a dramatic increase in the numbers and biomass of diatoms (Figure 2) in all size classes while other algal classes remained relatively static, with the exception of a small increase in *Phaeocystis* sp. [Hoffmann et al., 2006]. Towards the end of the experiment there was an apparently rapid export of organic matter from the mixed layer to deeper waters as evidenced by the presence of both the diatom marker fucoxanthin and low phaeopigment to chlorophyll ratios in samples collected throughout a water column of 4000 m [Peeken et al., 2005]. This sinking organic matter was in part comprised of particular diatom species like *Chaetoceros dictyota* whose biomass in the upper water column at this time decreased rapidly (Figure 2). The presence of particle aggregates below the mixed layer was also seen as spikes in the fluorescence and transmission (not shown) profiles at this time (Figure 1).

¹Auxiliary materials are available in the HTML. doi:10.1029/2006GL028080.

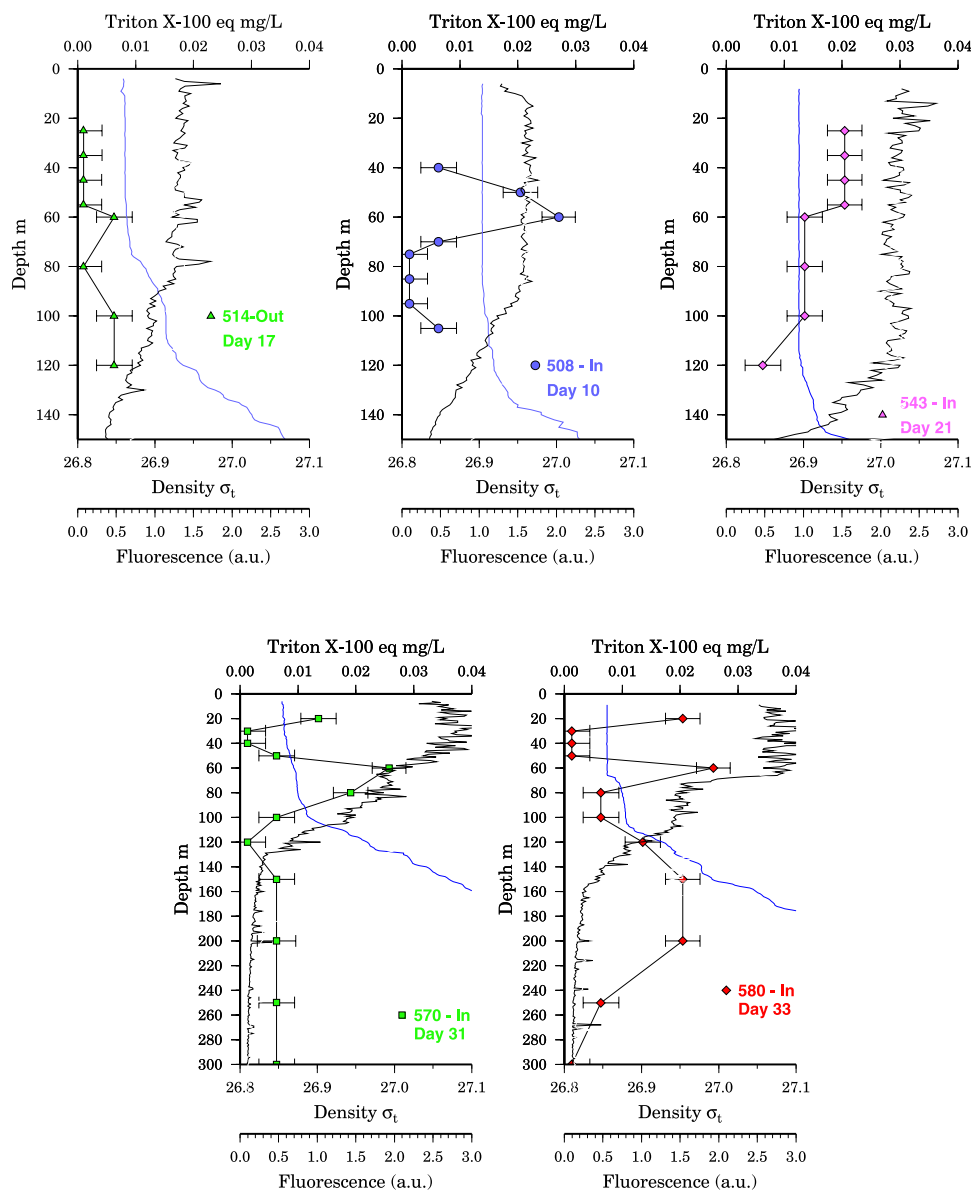


Figure 1. Temporal development of the vertical profile of SAS (symbols) in the water column during EIFEX. (top left) Station 514 (green triangles – OUT patch). (top middle) Stations 508 (blue circles – 10 days after 1st infusion). (top right) Station 543 (pink diamonds – 21 days after the 1st Fe infusion). (bottom left) Station 570 (green squares – 31 days after the 1st Fe infusion). (bottom right) Station 580 (red diamonds – 33 days after the 1st Fe infusion). Density profiles (solid blue lines) and fluorescence profiles (solid black lines) are also shown for each station to indicate the MLD and the distribution of phytoplankton at this time. Data from station 427 is not shown as the SAS concentrations were at or below the detection limit (0.001 mg L^{-1} Triton X-100 eq). Note the scale change in the depth for the last two stations.

[10] Over the course of EIFEX the upper water column inventory of SAS (0–100 m: calculated using the trapezoidal rule) of the Fe fertilized patch apparently increased over the first 20 days and then slightly decreased afterwards (Figure 2). The only out station sampled (day 17) still had low integrated SAS concentrations similar to the in patch initial conditions.

4.3. Enhanced SAS at Density Discontinuities

[11] Although SAS is thought to be produced by phytoplankton [Zutic *et al.*, 1981] the vertical distribution pattern did not reflect bulk fluorescence or chlorophyll profiles. A strong feature in the vertical profiles (Figure 1) is the

presence of zones of high SAS that are often aligned with small steps in the CTD derived density profiles. This phenomena is most notably for days 31–33 at which time the AML had decreased due to a slight warming (density decrease) of the near surface waters. This enrichment in SAS may arise from accumulation of phytoplankton source material at density discontinuities [MacIntyre *et al.*, 1995]. An additional important factor in the case of samples below 50 m from both day 31 and 33 is that this depth is now below the 1% light level (approximately 40 m) at this time. Thus phytoplankton in turbulent overturns below the AML would no longer be mixed above the 1% light level and thus maybe light limited; a situation for some diatom species that

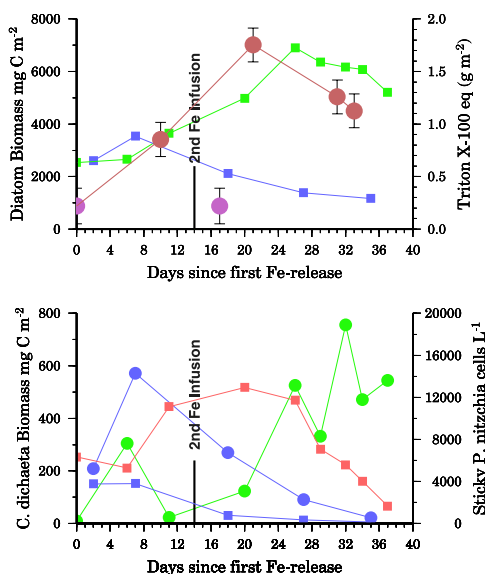


Figure 2. (top) Water column inventories (top 100 m) of the estimated total diatom biomass (squares: green, in patch; blue, out patch) and SAS (circles: red, in patch; purple, out patch), expressed as Triton X-100 equivalents, during EIFEX. The initial SAS inventory from station 427 is estimated as 3 times the detection limit (0.001 mg L^{-1} Triton X-100 eq) multiplied by the water column depth (100 m) giving rise to an estimate of $0.3 \pm 0.2 \text{ g m}^{-2}$ of Triton X-100 eq (95% confidence interval). (bottom) Average abundance (top 100 m) of ‘sticky’ *Pseudo-nitzschia* sp. (circles: green, in patch; blue, out patch) and the estimated biomass for the diatom *Chaetoceros dictyoceros* (squares: red, in patch; blue, out patch). Both were determined from direct microscopy counts.

leads to an increased release of organic exudates [Ignatiades and Fogg, 1973], which potentially may include SAS.

[12] The increase in SAS below the AML through the depth range 100–300 m is not consistent with the diffusion of a dissolved substance in water at this time. Typical diffusion rates for below the mixed layer in this region of the Antarctic [Cisewski et al., 2005] are $10^{-4} - 10^{-3} \text{ m}^2 \text{ s}^{-1}$ which gives a mass diffusion length ($\sqrt{4K_z t}$) of only 8.3–26.3 m over two days. Thus the SAS must be supplied by dissolution or exudation from sinking phytoplankton material. Even if the SAS were present as colloidal material in the samples they would need to be transported by particles to this depth as only Stokes settling of large organic particles (0.1–1 mm) would be fast enough to accomplish a descent of 100–200 m in 2 days or less. This implies then that the formation/release of SAS was very rapid during this time as part of the bloom began to sink out.

4.4. Possible Composition, Sources and Sinks of SAS

[13] SAS may be derived from a number of compounds including polysaccharides, proteins, humic acid and fulvic acid. Previous work using ONP detection of SAS has utilized changes in the height of the pre-peak of ONP to characterize the nature of the SAS present in the seawater [Gašparović and Čosović, 1995]. Using the same approach here it appeared that the initial SAS was strongly dominated by neutral polysaccharide components: consistent with results

from a phytoplankton bloom in the Adriatic [Gašparović et al., 1998a] and culture studies with the diatom *Phaeodactylum tricornerutum* [Gašparović et al., 1998b]. During day 31–33, when the EIFEX bloom was beginning to collapse, ONP pre-peak data suggested that while polysaccharides were still dominant, lipid and/or acidic polysaccharide material was also making a contribution to the SAS pool. Comparisons with measurements of total mono and polysaccharides and dissolved organic carbon (DOC) from EIFEX (S. Steigenberger, Role of TEP during EIFEX, manuscript in preparation, 2007) showed no apparent relationship with SAS, though this might be expected since surface active polysaccharides may contribute only a small fraction of the total polysaccharide or DOC pools.

[14] The most likely source, for SAS during EIFEX is from exudation products from diatoms which may be released under a variety of physiological stresses including nutrient/light limitation [Ignatiades and Fogg, 1973; Mopper et al., 1995; Waite et al., 1995]. Furthermore integrated SAS levels in EIFEX do appear to be closely related to the total diatom biomass in the upper water column (Figure 2). An alternative mechanism for SAS release is zooplankton grazing on diatoms [Kujawinski et al., 2002], we examined this further during EIFEX by undertaking a small scale grazing experiment with individual copepods, representative of the EIFEX patch [Jansen et al., 2006], collected close to the time of Station 580. Only *Calanus propinquus* showed a significant change with a tripling of SAS levels (0.006 to 0.020 mg L^{-1} Triton X-100 eq), while experiments with individual *Calanus simillimus* and *Pleuromamma* showed no change in SAS over 24 hours, as did control samples with no copepods present. No experiments were conducted with protozoans. Whether the results here can be attributable to species or simply to the physiological status of the individual zooplankton is beyond the scope of the present work.

[15] The sink terms for SAS are less clear however; the components of SAS may be utilized by bacteria, altered by photochemistry, adsorbed on particles or accumulated on rising bubbles which either transport it to the microlayer or form TEP in situ [Azetsu-Scott and Passow, 2004; Zhou et al., 1998]. From the limited data set we have here it is not possible to determine what the major loss term for SAS was in the water column.

4.5. SAS and Sticky Diatoms

[16] An interesting feature at the time of the bloom collapse was the apparent increase in “sticky” diatoms, as defined by the observation of attached cellular debris to diatom chains, most notably *Pseudo-nitzschia* sp (Figure 2) mostly as *P. lineola*, *P. turgidula*, and *P. turgiduloides*. Most of the material attached to these diatoms possibly originated from crustacean fecal pellets as suggested earlier in studies from the Pacific Ocean [Buck and Chavez, 1994]. The maximum in “sticky” diatoms occurs some day after the apparent maxima in integrated SAS in the water column (Figure 2) – though as Buck and Chavez noted previously it may take several days for cells to divide and the diatom chains to lengthen after initial attachment to the fecal pellet. Diatoms are also well known to produce sugar containing compounds on their cell surfaces which modify their stickiness and thus their aggregation behavior [Waite et al., 1995]. Additionally some

species of *Pseudo-nitzschia* are known to produce the toxin domoic acid as a low affinity uptake system for Fe and Cu [Wells et al., 2005], raising the possibility a significant part of the SAS activity may have been from compounds produced in response to increasing iron stress, light limitation, or as a grazing defense as the bloom progressed.

5. Conclusions

[17] The temporal changes of SAS during an open ocean iron induced phytoplankton bloom are reported for the first time. SAS levels increased during the course of the experiment and often appeared to be highest in regions of density discontinuities, presumably from material accumulation. This initial study highlights the need for more work focused on elucidating the molecules that make up SAS in seawater and their sources and sinks. Processes involving SAS in the sea surface microlayer may influence the air-sea exchange of climate important gases and aerosols and it is to the microlayer that future work should be concentrated; despite the technological challenge that must be faced for microlayer sampling under the sea conditions typically encountered in the Southern Ocean.

[18] **Acknowledgments.** The authors would like to show their deep thanks and appreciation to the crew of the R.V. *Polarstern*, for all their efforts in helping us throughout ANTXXI-3. Thanks also to the Chief Scientist, Victor Smetacek and to the AWI for making this cruise possible. This work was in part supported by a DFG grant awarded to P. Croot (CR145/4-1) and U. Passow (PA424/6-1). This work is a contribution to the German SOLAS (SOPRAN) program.

References

- Azetsu-Scott, K., and U. Passow (2004), Ascending marine particles: Significance of transparent exopolymer particles (TEP) in the upper ocean, *Limnol. Oceanogr.*, **49**, 741–748.
- Brainerd, K. E., and M. C. Gregg (1995), Surface mixed and mixing layer depths, *Deep Sea Res., Part I*, **42**, 1521–1543.
- Buck, K. R., and F. P. Chavez (1994), Diatom aggregates from the open ocean, *J. Plankton Res.*, **16**, 1449–1457.
- Cisewski, B., V. H. Strass, and H. Prandke (2005), Upper-ocean vertical mixing in the Antarctic polar front zone, *Deep Sea Res., Part II*, **52**, 1087–1108.
- Frew, N. M., J. C. Goldman, M. R. Dennett, and A. S. Johnson (1990), Impact of phytoplankton-generated surfactants on air-sea gas exchange, *J. Geophys. Res.*, **95**, 3337–3352.
- Gašparović, B., and B. Čosović (1994), Electrochemical estimation of the dominant type of surface-active substances in seawater samples using o-nitrophenol as a probe, *Mar. Chem.*, **46**, 179–188.
- Gašparović, B., and B. Čosović (1995), Electrochemical reduction of o-nitrophenol as a tool for the rough characterization of organic matter in seawater samples, *Electroanalysis*, **7**, 1136–1142.
- Gašparović, B., and B. Čosović (2001), Distribution of surface-active substances in the northern Adriatic Sea, *Mar. Chem.*, **75**, 301–313.
- Gašparović, B., V. Vojvodić, and B. Čosović (1997), Characterization of organic matter in fractionated seawater samples using o-nitrophenol as an electrochemical probe, *Anal. Chim. Acta*, **338**, 179–190.
- Gašparović, B., V. Vojvodić, and B. Čosović (1998a), Excretion of organic matter during an experimental phytoplankton bloom followed using o-nitrophenol as an electrochemical probe, *Croat. Chem. Acta*, **71**, 271–284.
- Gašparović, B., B. Čosović, and V. Vojvodić (1998b), Contribution of organic acids to the pool of surface active substances in model and marine samples using o-nitrophenol as an electrochemical probe, *Org. Geochem.*, **29**, 1025–1032.
- Gašparović, B., M. Plavšić, B. Čosović, and M. Reigstad (2005), Organic matter characterization and fate in the sub-arctic Norwegian fjords during the late spring/summer period, *Estuarine Coastal Shelf Sci.*, **62**, 95–107.
- Hardy, J. T., C. W. Apts, E. A. Creclius, and G. W. Fellingham (1985), The sea-surface microlayer: Fate and residence times of atmospheric metals, *Limnol. Oceanogr.*, **30**, 93–101.
- Hillebrand, H., C. D. Durselen, D. Kirschtel, U. Pollinger, and T. Zohary (1999), Biovolume calculation for pelagic and benthic microalgae, *J. Phycol.*, **35**, 403–424.
- Hoffmann, L., I. Peeken, K. Lochte, P. Assmy, and M. Veldhuis (2006), Different reactions of Southern Ocean phytoplankton size classes to iron fertilisation, *Limnol. Oceanogr.*, **51**, 1217–1229.
- Hunter, K. A. (1980), Processes affecting particulate trace metals in the sea surface microlayer, *Mar. Chem.*, **9**, 49–70.
- Hunter, K. A., and P. S. Liss (1981), Polarographic measurement of surface-active material in natural waters, *Water Res.*, **15**, 203–215.
- Ignatiades, L., and G. E. Fogg (1973), Studies on the factors affecting the release of organic matter by *Skeletonema costatum* (Greville) Cleve in culture, *J. Mar. Biol. Assoc. U. K.*, **53**, 937–956.
- Jansen, S., C. Klaas, S. Krägefsky, L. von Harbou, and U. Bathmann (2006), Reproductive response of the copepod *Rhincalanus gigas* to an iron-induced phytoplankton bloom in the Southern Ocean, *Polar Biol.*, **29**, 1039–1044.
- Kujawinski, E. B., J. W. Farrington, and J. W. Moffett (2002), Evidence for grazing-mediated production of dissolved surface-active material by marine protists, *Mar. Chem.*, **77**, 133–142.
- MacIntyre, S., A. L. Alldredge, and C. C. Gotschalk (1995), Accumulation of marine snow at density discontinuities in the water column, *Limnol. Oceanogr.*, **40**, 449–468.
- Menden-Deuer, S., and E. J. Lessard (2000), Carbon to volume relationships for dinoflagellates, diatoms, and other protist plankton, *Limnol. Oceanogr.*, **45**, 569–579.
- Mopper, K., J. A. Zhou, K. S. Ramana, U. Passow, H. G. Dam, and D. T. Drapeau (1995), The role of surface-active carbohydrates in the flocculation of a diatom bloom in a mesocosm, *Deep Sea Res., Part II*, **42**, 47–73.
- Passow, U. (2000), Formation of transparent exopolymer particles, TEP, from dissolved precursor material, *Mar. Ecol. Prog. Ser.*, **192**, 1–11.
- Passow, U., A. L. Alldredge, and B. E. Logan (1994), The role of particulate carbohydrate exudates in the flocculation of diatom blooms, *Deep Sea Res., Part I*, **41**, 335–357.
- Peeken, I., et al. (2005), Export of fresh algal material during the Southern Ocean iron fertilisation experiment, EIFEX, paper presented at Summer Meeting 2005, Am. Soc. of Limnol. and Oceanogr., Santiago de Compostela, Spain, 19–24 June.
- Röttgers, R., F. Colijn, and M. Dibbern (2005), Algal physiology and biooptics, *Ber. Polarforsch.* **500**, pp. 82–88, Alfred Wegener Inst. für Polar und Meeresforsch., Bremerhaven, Germany.
- Shine, J. P., and G. T. Wallace (1996), Flux of surface-active organic complexes of copper to the air-sea interface in coastal marine waters, *J. Geophys. Res.*, **101**, 12,017–12,026.
- Smetacek, V., et al. (2005), The expedition ANT XXI/3 of R/V *Polarstern*, *Ber. Polarforsch.* **500**, pp. 1–134, Alfred Wegener Inst. für Polar und Meeresforsch., Bremerhaven, Germany.
- Strass, V., B. Cisewski, S. Gonzalez, H. Leach, K.-D. Loquay, H. Prandke, H. Rohr, and M. Thomas (2005), The physical setting of the European Iron Fertilization Experiment 'EIFEX' in the Southern Ocean, *Ber. Polarforsch.* **500**, pp. 15–46, Alfred Wegener Inst. für Polar und Meeresforsch., Bremerhaven, Germany.
- Tsai, W., and K. Liu (2003), An assessment of the effect of sea surface surfactant on global atmosphere-ocean CO₂ flux, *J. Geophys. Res.*, **108**(C4), 3127, doi:10.1029/2000JC000740.
- Utermöhl, H. (1958), Zur Vervollkommnung der quantitativen Phytoplankton-Methodik, *Mitt. Int. Ver. Theor. Angew. Limnol.*, **9**, 1–38.
- Waite, A. M., R. J. Olson, H. G. Dan, and U. Passow (1995), Sugar-containing compounds on the cell surfaces of marine diatoms measured using concanavalin A and flow cytometry, *J. Phycol.*, **31**, 925–933.
- Wallace, G. T. (1982), The association of copper, mercury and lead with surface-active organic matter in coastal seawater, *Mar. Chem.*, **11**, 379–394.
- Wells, M. L., C. G. Trick, W. P. Cochlan, M. P. Hughes, and V. L. Trainer (2005), Domoic acid: The synergy of iron, copper, and the toxicity of diatoms, *Limnol. Oceanogr.*, **50**, 1908–1917.
- Zhou, J., K. Mopper, and U. Passow (1998), The role of surface-active carbohydrates in the formation of transparent exopolymer particles by bubble adsorption of seawater, *Limnol. Oceanogr.*, **43**, 1860–1871.
- Zutic, V., B. Čosović, E. Marcenko, and N. Bihari (1981), Surfactant production by marine phytoplankton, *Mar. Chem.*, **10**, 505–520.

P. Assmy, S. Jansen, U. Passow, and V. H. Strass, Alfred Wegener Institute for Polar and Marine Research, Am Handelshafen 12, D-27570 Bremerhaven, Germany.

P. L. Croot, Forschungsbereich Marine Biogeochemie, Leibniz-Institut für Meereswissenschaften at University of Kiel, Dienstgebäude Westufer, Düsternbrooker Weg 20, D-24105 Kiel, Germany. (pcroot@ifm-geomar.de)