# **Science Advances NAAAS**

# Supplementary Materials for

## **Caribbean salinity anomalies contributed to variable North Atlantic circulation and climate during the Common Era**

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## **The PDF file includes:**

Supplementary Text Figs. S1 to S7 Tables S1 to S4 References

## **Other Supplementary Material for this manuscript includes the following:**

Data S1 and S2

## **Supplementary Text**

### **Comparison with the nearby core-top SST estimates**

As can be seen from table S2, our SST estimate for the M35003-6 core-top is ~1.3±0.5 °C lower than previously obtained for this location (53) and for a core M35/1-2-1 that is 10 km apart (21). While these previous studies examined only larger specimens (>250 μm; ref. 21,53), we analyzed a large variety of *G. ruber* (white) shell sizes (>150 μm), including also small individuals that tend to prevail during the winter season (61). Therefore, the slightly colder core-top SST obtained here can be the result of either different calcification rates or seasonal preferences. Despite some discrepancies, all core-top SST and SSS estimates are consistent with the mean annual temperature and salinity of the upper water column in the Tobago Basin (Fig. 3 in the main text).

### **Age model**

Dating of Caribbean cores is currently associated with centennial-scale age errors, due to uncertainties in both radiocarbon calibration and local marine reservoir correction  $(ΔR)$  values (62,63). Additional errors in the range of decades to centuries may result from the  $\Delta R$  variability associated with changes in Orinoco and Amazon discharges (63) and bioturbation. However, accurate age control in our sedimentary archive is essential to investigate the covariance (or lack thereof) between the southeastern Caribbean and Gulf of Mexico SSS.

The initial age model for core M35003-6 was based on four <sup>14</sup>C accelerator mass spectrometer (AMS) dates using *G. ruber* (white). The dating was carried out at Leibniz Laboratory for Radiometric Dating and Stable Isotope Research, Kiel, Germany (fig. S2; table S3). The AMS 14C

ages were calibrated using *RBacon* (64) and a recent Marine20 calibration curve with a modern global ocean age increased by  $~150$  years when compared to 400 years previously used in Marine13 (62). The  $\Delta R$  value was taken from the  $\Delta R$  database updated for Marine20 (http://calib.org/marine) by averaging values from 7 nearest points in the southeastern Caribbean region (-248±144 years, Trinidad, Tobago, Los Testigos; refs. 63,65). The high positive value from Grenada (282±24 years, ref. 63) was not included in our calculation, because samples associated with the Atlantic influence tend to have strongly negative values (63,65,66). In addition, a modern age for the core-top was assumed because our box core contained an intact and excellently preserved sediment-water interface. We note, however, that the resulting age model suggests constant sedimentation rates over the past 2,000 years (fig. S2), which is not supported by our sedimentological data. In particular, high number of coarse lithic grains  $(>150 \mu m)$  at the base of our box core (lower ~10 cm) indicates lowered deposition of fine-grained fluvial material (fig. S2), and, therefore, a pronounced change in sedimentation rates is expected, in agreement with earlier reconstructions at this site (e.g., ref. 20).

To account for the unavoidable dating ambiguity resulting from both radiocarbon-date uncertainties and variable sedimentation rates, we build our chronology on alignment to the Gulf of Mexico SST record (core 2010-GB2-MCA, ref. 14). We updated the original age model for core 2010-GB2-MCA using the Marine20  $\Delta R$  database ( $\Delta R$ =-164±9 years, ref. 66) and a Marine20 calibration curve (62). Paired Mg/Ca- $\delta^{18}O$  values from this core were reprocessed in PSU Solver (23) using the same calibrations as proposed in the original publication (14). Analytical and sampling uncertainties were incorporated into overall uncertainty derived in PSU Solver (for Mg/Ca, combined 2 $\sigma$  analytical and sampling uncertainty of  $\pm 0.10$  mmol/mol; for  $\delta^{18}O$ , only 2 $\sigma$ analytical uncertainty of  $\pm 0.12$  ‰). Each sampling depth was also associated with a  $2\sigma$  age uncertainty of  $\pm 150$  years based on radiocarbon dating (table S4). The estimated average  $1\sigma(2\sigma)$ uncertainty is  $\pm 0.6$  (1.1) °C for SST and  $\pm 0.3$  (0.6) units for SSS.

To align Caribbean and Gulf of Mexico SST records, we make use of two climate events that are distinct in both records: 1) the pronounced SST cooling prior to the Medieval Peak Warmth and 2) the warmer spell of the LIA (fig. S2). Our assumption of synchronous SST changes across these two basins is based on observations (67) and correlation analysis (ref. 14, their fig. S3). By using the Gulf of Mexico SST record as a reference we assume that uncertainties in core 2010-GB2- MCA chronology are lower, which is justified e.g., by smaller errors in  $\Delta R$  (66) and/or lower bioturbation effects. The latter is supported by radionuclide data showing that active sediment mixing on the seafloor in the oligotrophic Gulf of Mexico is limited to upper 4 cm (68). In contrast, the >8 cm thickness of bioturbated sediment layer in the southeastern Caribbean cannot be ruled out, due to its relatively high primary productivity and organic carbon flux under riverine fertilization (27,69). The resultant sedimentation rates over the last 1,700 years between  $\sim$ 10 and 45 cm kyr<sup>-1</sup> are consistent with earlier reconstructions on the Holocene at this site  $(20,70)$ . However, we suggest that the core section older than 550 C.E. is characterized by low sedimentation rates, based on high number of coarse grains indicative of reduced fluvial inputs. Due to the lack of adequate age control, we do not interpret this older part of the record.



**Fig. S1.** Proxy data from core M35003-6 over the past 1,700 years. **(A, B)**  $\delta^{18}O$  and Mg/Ca values in *G. ruber* (white) with their  $2\sigma$  analytical and sampling uncertainties (see Materials and Methods). **(C, D)** Mg/Ca- $\delta^{18}O$ -based SST and SSS records;  $\pm 1\sigma$  uncertainties calculated in PSU Solver (see Materials and Methods). Original and smoothed records (3-point moving average) are



**Fig. S2. Age model for core M35003-6.** Smoothed (3-point moving average) SST record from core M35003-6 is plotted against timescales derived from **(A)** radiocarbon dates and **(C)** core-tocore correlations (i.e., the final age model). Dashed black lines show the alignment of M35003-6 SST record to  $(B)$  core 2010-GB2-MCA SST, smoothed using  $2\sigma$  age model uncertainty in PSU Solver (14). Red triangles and orange crosses show calibrated radiocarbon dates  $(\pm 2\sigma)$  for M35003-6 and 2010-GB2-MCA, respectively. Also shown is the abundance of lithic grains in the fraction >150 µm **(D)** and sedimentation rates in core M35003-6 **(E)**. Grey arrow in **(D)** denotes an abrupt reduction in the abundance of coarse lithic grains at c. 550 C.E. that must have coincided with an increase in fluvial inputs and, thus, higher sedimentation rates (see supplementary text).



**Fig. S3. Proportions of selected planktic foraminifera species in core M35003-6 compared to reconstructed sea-surface salinities (original and 3-point moving average smoothed records).** For assemblage data, grey lines show 95 % confidence bounds for smoothed records (see Materials and Methods). Blue bars denote centennial cooling/salinification events in the southeastern Caribbean. LALIA – Late Antique Little Ice Age, LIA-Little Ice Age, MWP-Medieval Warm Period.



**Fig. S4. Scree plot of the eigenvalues for the 10 principal components (PCs, blue line) of the Principal Component Analysis (PCA) performed on the planktic foraminifera assemblage dataset from core M35003-6**. Error bars indicate the 95 % confidence interval computed by bootstrapping ( $n = 1000$ ). PCs 1, 2, and 3 are considered significant, as the eigenvalues for these three PCs lie above the eigenvalues expected under a random model ("broken stick", red dashed line). We note, however, that the broken stick is inside the 95 % confidence interval for the fourth PC. PCA was performed using PAST software (48).



**Fig. S5. Comparison between reconstructed salinities and scores of the principal components (PCs) of the planktic foraminifera assemblage dataset in core M35003-6.** PC1 (orange line) explains 33 % of the dataset variance and shows high positive loadings for the proportions of *Globigerinoides sacculifer* and *Neogloboquadrina dutertrei* (0.43 and 0.34, respectively) and high negative coefficients for *Globigerinoides ruber* (pink) and *Globigerinoides ruber* (white) (-0.66 and -0.41, respectively). PC2 (red line) accounts for 27 % of the variance and represents a gradient between *G. ruber* (pink) (0.56) and *G. ruber* (white) (-0.63). PC3 (blue line, 17 % of the variance) represents a gradient where positive values represent the combined proportions of *G. ruber* (white) (0.42) and *N. dutertrei* (0.34) and negative values are associated with *G. glutinata* (-0.72). Correlations between the paleosalinity record and the scores of PCs 1-3 are insignificant ( $p>0.05$ ). Blue bars denote centennial cooling/salinification events in the southeastern Caribbean. LALIA – Late Antique Little Ice Age, LIA-Little Ice Age, MWP-Medieval Warm Period.



**Fig. S6. Caribbean – Gulf of Mexico salinity seesaw on radiocarbon-derived timescales.**  Comparison of SST and SSS reconstructions from the Caribbean core M35003-6 (black, this study) and Gulf of Mexico core 2010-GB2-MCA (purple, ref. 14), plotted on the radiocarbonbased timescales. Black triangles and purple crosses show calibrated radiocarbon dates  $(\pm 2\sigma)$  for M35003-6 and 2010-GB2-MCA, respectively. See supplementary text for details on the age model construction. For radiocarbon dates see tables S3 and S4. Grey bars denote centennial cooling/salinification events in the southeastern Caribbean as discussed in the main text.



**Fig. S7. BEAST decomposition of the sea surface salinity (SSS) reconstructed from core M35003-6.** Smoothed (3-point moving average) SSS timeseries (black line) were decomposed into abrupt changes (green bars) and a fitted trend (orange line). The median number of changepoints in the SSS trend is five (12 % probability; dashed grey lines). Green bars show probability distribution of having a changepoint in the trend at each point of time. Higher peaks indicate a higher chance of being a changepoint only at that particular single point in time and does not necessarily mean a higher chance of observing a changepoint around that time. The summed probability for each change point is indicated with numbers. BEAST, a Bayesian change point detection algorithm (24), supports statistical significance of SSS changes between 550 and 900 C.E. and an increase in surface salinity at 1550 C.E.

| # in<br>Fig. 1          | <b>Material</b>                                                             | Location                                                    | <b>Coordinates</b>                                                                   | Water depth                            | Reference   |
|-------------------------|-----------------------------------------------------------------------------|-------------------------------------------------------------|--------------------------------------------------------------------------------------|----------------------------------------|-------------|
| $\mathbf{1}$            | Box core M35003-6                                                           | Tobago Basin,<br>southeastern<br>Caribbean                  | 12°05'06"N<br>61°14'42"W                                                             | 1299 m                                 | this study  |
| $\boldsymbol{2}$        | Lacustrine sediment<br>core                                                 | South Sawtooth<br>Lake, Ellesmere<br>Island, Canada         | 79°21'03"N<br>83°57'07"W                                                             | $80 \text{ m}$<br>(lake)               | Ref. 15     |
| $\overline{\mathbf{3}}$ | Arctica islandica site                                                      | north Icelandic shelf                                       | 66°31'35"N<br>18°11'44"W                                                             | 80 <sub>m</sub>                        | Refs. 12,13 |
| $\overline{\mathbf{4}}$ | RAPiD-35-COM<br>(box core RAPiD-35-<br>25B and piston core<br>RAPiD-35-14P) | Eirik Drift, eastern<br>Labrador Sea                        | Box core-<br>57°30'28"N<br>48°43'24"W<br>Piston core-<br>57°30'15" N<br>48°43'20" W  | ~1.3485~m                              | Ref. 10     |
| 5                       | RAPiD-21-COM<br>(box core RAPiD-21-<br>12B and kasten core<br>RAPiD-21-3K)  | Gardar Drift,<br>subpolar North<br>Atlantic                 | Box core $-$<br>57°27'05"N<br>27°54'14"W<br>Kasten core-<br>57°27'05"N<br>27°54'32"W | 2630 m                                 | Ref. 10     |
| 6                       | <b>KNR-178-48JPC</b>                                                        | off Cape Hatteras,<br>western North<br>Atlantic             | 35°46'N<br>74°27'W                                                                   | 2009 m                                 | Ref. 47     |
| $\overline{7}$          | 2010-GB2-MCA                                                                | Garrison Basin,<br>northern Gulf of<br>Mexico               | 26°40'11″N<br>93°55'13"W                                                             | 1776 m                                 | Ref. 14     |
| 8                       | Cuba Grande (CG)<br>stalagmite                                              | Dos Anos Cave,<br>northwestern Cuba                         | 22°23'N<br>83°58'W                                                                   | 100 m above<br>sea level (cave)        | Ref. 38     |
| $\boldsymbol{9}$        | PDR-1 stalagmite                                                            | Perdida Cave,<br>central Puerto Rico                        | 18°N<br>$67^{\circ}$ W                                                               | 350-400 m<br>above sea level<br>(cave) | Ref. 37     |
| 10                      | Pond sediment core<br>(GT3)                                                 | Grape Tree Pond,<br>Jamaica                                 | 17°53'37"N<br>76°37'06"W                                                             | $1 \text{ m} \text{ (pond)}$           | Ref. 36     |
| 11                      | Lacustrine sediment<br>core                                                 | Lake Punta Laguna,<br>Yucatan Peninsula,<br>southern Mexico | $20^{\circ}38'N$<br>87°37'W                                                          | $6 \text{ m}$ (lake)                   | Ref. 33     |
| 12                      | YOK-I stalagmite                                                            | Yok Balum Cave,<br>Yucatan Peninsula<br><b>Belize</b>       | 16°12'31"N<br>89°04'24"W                                                             | 366 m above<br>sea level (cave)        | Ref. 35     |

**Table S1. List of proxy records used in the study.**

| Core                                   | M35003-6 <sup>a</sup> | $M35003-6b$      | $M35/1-2-1$ <sup>c</sup> |
|----------------------------------------|-----------------------|------------------|--------------------------|
|                                        |                       |                  |                          |
| Sample depth (cm)                      | $0 - 0.5$             | $0-1$            | $0-1$                    |
| G. ruber (white)                       | sensu stricto         | n/a              | sensu lato               |
| Sediment fraction                      | $>150 \mu m$          | 355-400 um       | $250 - 313 \mu m$        |
| $Mg/Ca$ (mmol/mol) $\pm 2\sigma$       | $4.41 \pm 0.11$       | $4.72 \pm 0.15$  | $4.91 \pm 0.11$          |
| $\delta^{18}O(%) \pm 2\sigma$          | $-1.83 \pm 0.14$      | $-1.69 \pm 0.14$ | $-1.93 \pm 0.16$         |
| SST $({}^{\circ}C)^{d} \pm 1\sigma$    | $26.70 \pm 0.33$      | $27.99 \pm 0.33$ | $28.02 \pm 0.46$         |
| SSS (units) <sup>d</sup> $\pm 1\sigma$ | $35.57 \pm 0.21$      | $35.82 \pm 0.23$ | $35.84 \pm 0.28$         |

**Table S2. SST and SSS estimates for the Tobago Basin core-tops.**

<sup>a</sup> Mg/Ca and  $\delta^{18}$ O values from this study

 $<sup>b</sup>$  Mg/Ca values from ref. 53,  $\delta^{18}$ O values from ref. 71</sup>

 $\rm^c$  Mg/Ca and  $\delta^{18}$ O values from ref. 21

<sup>d</sup> Based on paired Mg/Ca- $\delta^{18}O$  using PSU Solver (ref. 23, see Materials and Methods)

**Table S3. Tie Points and Radiocarbon Age Control for core M35003-6.**

| Mid-<br>Depth<br>(cm) | Tie point<br>(Yr. C.E.) | <b>Lab Code</b><br>for ${}^{14}C$ date | Conv. Age<br>(Yr B.P.) | <b>Median Probability</b><br>Calendar Age*<br>(Yr C.E.) | <b>Calendar Age Range</b><br>$(2\sigma, Yr C.E.B.C.E.)$ |
|-----------------------|-------------------------|----------------------------------------|------------------------|---------------------------------------------------------|---------------------------------------------------------|
| 0.25                  | 1996                    |                                        |                        |                                                         |                                                         |
| 4.14                  | 1853                    |                                        |                        |                                                         |                                                         |
| 6.28                  | 1612                    |                                        |                        |                                                         |                                                         |
| 22.82                 | 889                     |                                        |                        |                                                         |                                                         |
| 28.45                 | 749                     |                                        |                        |                                                         |                                                         |
| 2.75                  |                         | KIA $50748**$                          | $560 \pm 20$           | 1853                                                    | 1675-1981                                               |
| 10.25                 |                         | KIA 50749**                            | 795±20                 | 1505                                                    | 1225-1788                                               |
| 26.25                 |                         | KIA 50750**                            | $1440 \pm 20$          | 852                                                     | 519-1151                                                |
| 44.75                 |                         | KIA 50751**                            | $2200\pm25$            | 31                                                      | $-382 - 420$                                            |

\*Radiocarbon dates for all depths were calibrated in BACON (64).

\*\*Radiocarbon dates were not used in the final age model, which was built on correlation to core 2010-GB2-MCA (4 tie points). In addition, modern age (1996 C.E.) was assumed for the core-top of M35003-6 (see supplementary text, fig. S2).

**Table S4. Radiocarbon Age Control for core 2010-GB2-MCA.**

| Mid-<br>Depth<br>(cm) | Conv. Age<br>(Yr B.P.)           | <b>BACON</b> Median<br><b>Probability Calendar</b><br>Age ( $Yr$ C.E.) | Calendar Age Range (2σ,<br>$Yr$ C.E./B.C.E.) |
|-----------------------|----------------------------------|------------------------------------------------------------------------|----------------------------------------------|
| 0.25                  | $>$ Modern $\pm$ 30 <sup>*</sup> | 1998                                                                   | 1936-2059                                    |
| 9.65                  | $855 \pm 30$                     | 1475                                                                   | 1326-1642                                    |
| 19.25                 | $1450 \pm 30$                    | 934                                                                    | 770-1105                                     |
| 34.75                 | $2405 \pm 30$                    | $-124$                                                                 | $-325-101$                                   |

\*Radiocarbon dates for all depths (from ref. 14) were re-calibrated in BACON (64) using the Marine20 ΔR database (ΔR=-164±9 years, ref. 66) and a Marine20 calibration curve (62), except for the core-top, for which a modern age was assumed.

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