

Text S1. Variations of *Pn* propagation in St. Paul and Romanche transforms

 St. Paul system and Romanche are two transform faults of the equatorial Atlantic, with different morphological and geodynamic features. St. Paul system is formed by four transform faults (Transform *A, B, C, D*) offsetting ~630 km between the two Mid-Atlantic Ridge (MAR) axes (Maia et al. 2016), as seen in Fig. S4. The Romanche transform is another extensive and complex composed of two transform valleys: one that extends over the 900 km offset between the two MARs and the second located in a further north area, close to the east side ridge. St. Paul has different deformation styles along the four transform faults: Transform *A* contains a 200-km long and 30-km wide submarine stepover with a shear zone and local thrust faults uplifting rocks 1.5 mm/year above sea level in St. Peter and St Paul islets (Campos et al. 2010; Maia et al. 52 2016), while Transform *C* is composed of corrugated oceanic core complex features where detachment faults expose lower crust and mantle rocks (Vincent et al. 2023). Some authors suggested the presence of a different deformation between the two transform valleys of Romanche, with a possible more vertical motion in the north-side valley (Bonati 1978; Ligi et al. 2002). In parallel, Abercrombie & Ekström (2001) used teleseismic signals to analyze focal depths of earthquakes in Romanche and identified a difference between the earthquake depths comparing the east and west sides of the Romanche transform fault, the east side having deeper focuses. Other papers provide new contrasting features regarding the low-temperature mantle in the northern area of the Romanche (Bonatti et al. 1996; Maia et al. 2020), thick and serpentinized crust (Gregory et al. 2021), containing micro seismicity in that area with deformation reaching 34 km (Yu et al. 2021).

 We plotted the magnitude adjustments of the strike-slip earthquakes along the St. Paul Romanche transform fault over the global lithospheric thickness modeled from Vs global tomography CAM2016 (Priestley et al. 2018, Fig.S3). In the St. Paul system, we observe a lithospheric thickness variation of ~8 km along the four transform faults. Most magnitude adjustments are positive in Transform *A*, with some events presenting negative adjustments close to the west MAR axis or over the thrust faults, reflecting different contrasts in local lithospheric stress in that transform fault. Differently, most earthquake magnitude adjustments were negative in Transform **C**, with the existence of a reduced lithosphere thickness caused by the low-magma activity in oceanic core complexes close to the ridge axis (Vincent et al. 2023). As an effect, there is a difference in the Pn wave velocities (de Melo et al. 2021) and propagation in that area

 when compared to Transform *A*. In Romanche, the lithospheric thickness varies ~10 km between the two MAR axes. We observed that most negative magnitude corrections on the east side were located in the thinner zone of the CAM2016 model, with positive station corrections in the thicker lithospheric side. A possible explanation is the difference in the upwelling mantle melting between the east and west zones of the Romanche transform, collaborating with the existence of a variation in the Pn propagation along the Romanche transform with a complex lithospheric structure between the two MARs.

82 **Tables.**

83

Table S1. List of Earthquakes Analyzed*

85

2014 Mar 11 17:42:01 0.59 -27.02 13 5.09 5.36 13 S 0.367 SPS

 87 * h= focal depth in km, N=number of observations; F = focal mechanism (S=strike-slip, N=Normal

88 faulting).

 $\frac{1}{2}$ Region: SPS = St. Paul system; RO-SPS = Romanche-St. Paul system; SPS-S= St. Paul system-

90 Strakhov; DSS = Doldrums system – south; DSN = Doldrums system – north; 15-20 = Fifteen-Twenty, 4-

91 $5^\circ N = 4-5^\circ N$; Chain-RO = Chain-Romanche; NA-SA = NA-SA plate boundary.

Table S2. Inversion Results* 93

94

95

96 $*$ Nobs = number of *Pn* amplitude measurements, Nsta = number of stations used, Neve =

97 number of events analyzed, $b = slope$ of amplitude-distance curve, $K = constant$ relating mb(Pn)

98 to Mw, plus attenuation beneath station, Region = applicable regions.

Pn, Wood-Anderson Peak Amplitudes, 2.74 Hz, (0.5-15 Hz), West

 Figure S1. Attenuation curve for the Western part of the equatorial MAR, with 1,663 observed *Pn* peak amplitudes in nanometers. The logarithm of peak amplitude normalized by subtracting the event moment magnitude Mw is plotted with *gray circles*. The calculated peak amplitudes in the least-squares inversion – after making station corrections and event magnitude adjustments (EMAs) – are plotted by *red circles* for comparison. The thick solid line is the amplitude- distance curve of *Pn* waves in the Western part of the equatorial Atlantic region with parameters as in Table S1.

Pn, Wood-Anderson Peak Amplitudes, 2.74 Hz, (0.5-15 Hz), East

 Figure S2. Attenuation curve for the Eastern part of the equatorial MAR, with 378 observed *Pn* peak amplitudes in nanometers. The logarithm of peak amplitude normalized by subtracting the event moment magnitude Mw is plotted with *gray circles*. The calculated peak amplitudes in the least-squares inversion – after making station corrections and event magnitude adjustments (EMAs) – are plotted by *red circles* for comparison. The thick solid line is the amplitude-distance curve of *Pn* waves in the Eastern part of the equatorial Atlantic region with parameters

as in Table S1.

123 Figure S3. Comparisons of station corrections and Vs distribution at a 150 km depth, and station corrections vs lithosphere-asthenosphere boundary (LAB) depths. a) Map with station corrections and Vs anomalies at 150km depth from SAAM23 model (Ciardelli et al. 2022), b) same with LAB (Lithosphere/Asthenosphere boundary) depth from CAM2016 model (Priestley *et al*. 2018). In the maps, positive and negative corrections are indicated by green and white triangles, respectively. c) station corrections are plotted against Vs distribution. Linear regression yields a correlation coefficient of 0.60 with a standard deviation of 0.20 magnitude units, and d) station corrections are plotted against LAB depths. Linear regression yields a correlation coefficient of 0.53 with a standard deviation of 0.22 magnitude units.

 Figure S4. Correlation of the event magnitude adjustments with the lithospheric thickness in the St. Paul system and Romanche transform fault. (upper left panel) Event magnitude adjustments for all strike-slip earthquakes in the St. Paul system (bottom left panel) Event magnitude adjustments of earthquakes in Romanche transform fault. The red and blue circles represent the negative and positive magnitude adjustments, respectively. (upper right panel) correlation of the event magnitude adjustments (circles) with the lithospheric thickness of CAM2016 in the St. Paul system, (bottom right panel) correlation of the event magnitude adjustments (circles) with the lithospheric thickness of CAM2016 in the Romanche transform fault. Background shade is coded by the longitude coordinate of the epicenter location in the fault.

147 **Fig. S5**. Comparison of the V_S distribution at 150 km depth and the amplitude-distance curves for Western (towards Brazil) and Eastern (toward African) paths. S-wave anomaly map at 150 km depth from the tomography model SAAM23 (Ciardelli *et al*. 2022). The thicker violet and gray lines indicate the approximate limits of the paths from all earthquakes to the Brazilian and African coasts, respectively. Double arrows "W" and "E" show the average limits of each region. The Vs anomalies in the Western side (towards Brazil), on average, are not as low as compared with the Eastern side (towards Africa). This is consistent with the higher amplitudes observed in the amplitude-distance curves for the Western side of the equatorial Mid-Atlantic Ridge, compared with the curve for the Eastern paths (African side) shown in Figure 7 in the main text. The gray area in Africa is outside the minimum resolution of Model SAAM23 (Ciardelli *et al*. 2022).

 Figure S6. Comparison of S-wave velocity distribution at 150 km depth from the tomography model SAM2019 (Celli *et al*. 2020) and *Pn* amplitude attenuation in the eastern and western 163 sides of the equatorial MAR. Lines as in Fig. S5. Vs is lower in the eastern side of the equatorial MAR, on average, compared to that of the western side. Lower *S*-wave velocities indicate stronger asthenosphere and probably higher attenuation. Figure based on Celli *et al.* (2020).

170 Figure S7. Comparisons of $m_b(Pn)$ and M_W values for 189 earthquakes. (a) $m_b(Pn)$ determined

from the peak amplitude compensated only with the amplitude-distance curve. The regression

line with a slope of 0.88 (*center solid line*) and lines of one orthogonal standard deviation of 0.20

magnitude units (*dashed lines*) are drawn. (b) *m*b(*Pn*) after making our event magnitude

174 adjustment (EMA) and station corrections. The regression yields a slope of 0.95 with a standard

deviation of only 0.06 magnitude units, indicating that an accurate approximation to the moment

176 magnitude in the range from Mw 4.95 to 7.14 is given by our values of $m_b(Pn)$ once the

magnitude adjustments (station and event corrections) are applied.

 Figure S8. Comparisons of $m_b(Pn)$ and M_W values for 189 earthquakes to examine the effect of 181 station corrections. (a) $m_b(Pn)$ determined by using event magnitude adjustment (EMA) but no station corrections. The regression yields a slope of 0.93 with a standard deviation of 0.10 magnitude units. Earthquake source types are indicated: normal (*circles*) and strike-slip (*red squares*). (b) *m*b(*Pn*) determined from the peak amplitude compensated with the amplitude- distance curve, event magnitude adjustment, and station corrections (SC). The regression line with a slope of 0.95 (*center solid line*) and lines of one orthogonal standard deviation of only 0.06 magnitude units (*dashed lines*) are drawn, indicating a reasonably weak effect of station 188 corrections on the values of $m_b(Pn)$.

References.

- Abercrombie, R. E., & Ekström, G., 2001. Earthquake slip on oceanic transform faults, *Nature*, **410,** 74-77.
- Bonatti, E., 1978. Vertical tectonism in oceanic fracture zones, *Earth and Planetary Science Letters*, **37**, 369-379.
- Bonatti, E., Ligi, M., Carrara, G., Gasperini, L., Turko, N., Perfiliev, S., & Sciuto, P. F., 1996. Diffuse impact of the Mid‐Atlantic Ridge with the Romanche transform: An ultracold ridge‐transform intersection, *J. Geophys. Res*., **101**, 8043-8054.
- Campos, T. F., Bezerra, F. H., Srivastava, N. K., Vieira, M. M., & Vita-Finzi, C., 2010. Holocene tectonic uplift of the St Peter and St Paul Rocks (Equatorial Atlantic) consistent with emplacement by extrusion, *Marine Geology*, **271**, 177-186.
- 201 de Melo, G. W., Parnell-Turner, R., Dziak, R. P., Smith, D. K., Maia, M., do Nascimento, A. F., 202 & Royer, J. Y., 2021. Uppermost mantle velocity beneath the Mid-Atlantic Ridge and transform faults in the Equatorial Atlantic Ocean, *Bull. Seismol. Soc. Am*, **111**, 1067- 1079.
- Gregory, E. P., Singh, S. C., Marjanović, M., & Wang, Z., 2021. Serpentinized peridotite versus thick mafic crust at the Romanche oceanic transform fault, *Geology*, **49**, 1132-1136.
- Ligi, M., Bonatti, E., Gasperini, L., & Poliakov, A. N., 2002. Oceanic broad multifault transform plate boundaries, *Geology*, **30**, 11-14.
- Maia, M., Sichel, S., Briais, A., Brunelli, D., Ligi, M., Ferreira, N., & Oliveira, P., 2016. Extreme mantle uplift and exhumation along a transpressive transform fault, *Nature Geoscience*, **9**, 619-623.
- Maia, M., & Brunelli, D., 2020. The Eastern Romanche ridge-transform intersection (Equatorial Atlantic): slow spreading under extreme low mantle temperatures. Preliminary results of the SMARTIES cruise, In *EGU General Assembly Conference Abstracts* (p. 10314).
- Vincent, C., Maia, M., Briais, A., Brunelli, D., Ligi, M., & Sichel, S., 2023. Evolution of a Cold Intra‐Transform Ridge Segment Through Oceanic Core Complex Splitting and Mantle
- Exhumation, St. Paul Transform System, Equatorial Atlantic, *Geochemistry, Geophysics,*
- *Geosystems*, **24**, e2023GC010870.
- Yu, Z., Singh, S. C., Gregory, E. P., Maia, M., Wang, Z., & Brunelli, D., 2021. Semibrittle
- seismic deformation in high-temperature mantle mylonite shear zone along the
- Romanche transform fault, *Science Advances*, **7**, eabf3388.