

Diapycnal Mixing in the Equatorial Cold Tongue and its Parameterization

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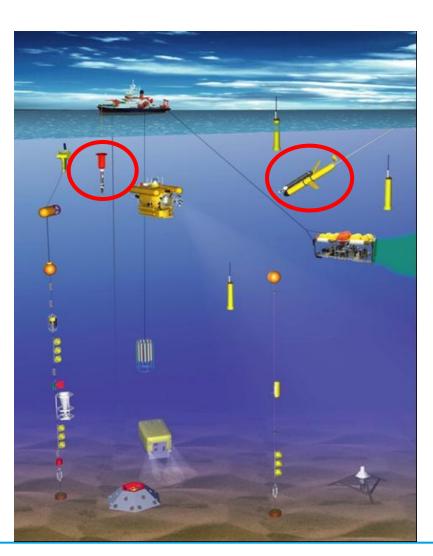


Mixing Observatories at GEOMAR



Ship-based microstructure systems



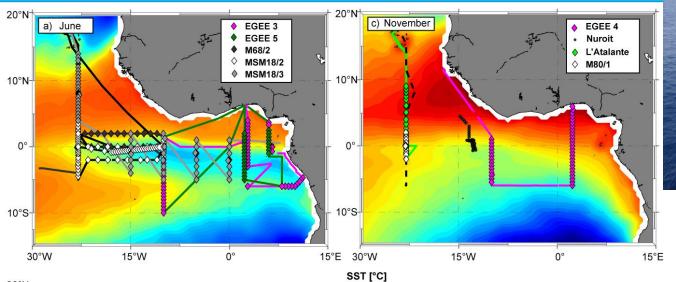


Autonomous microstructure platforms (MicroRider / Glider)

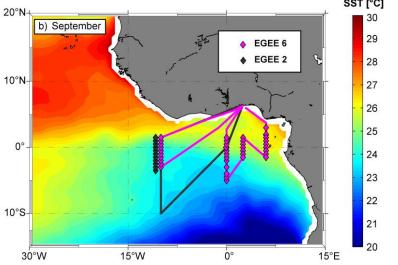




Ship-board microstructure measurements (2005-2011)





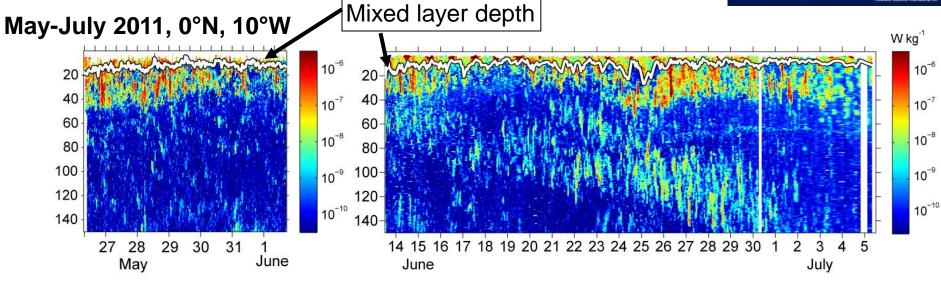


- repetitive microstructure sections within the cold tongue region from 11 cruises during different seasons
- individual stations with at least 3 profiles (>2000 profiles)
- shipboard ADCP measurements

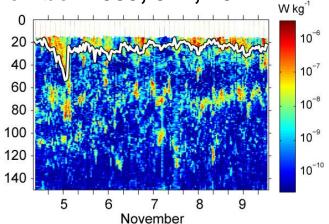


Time series of turbulent kinetic energy



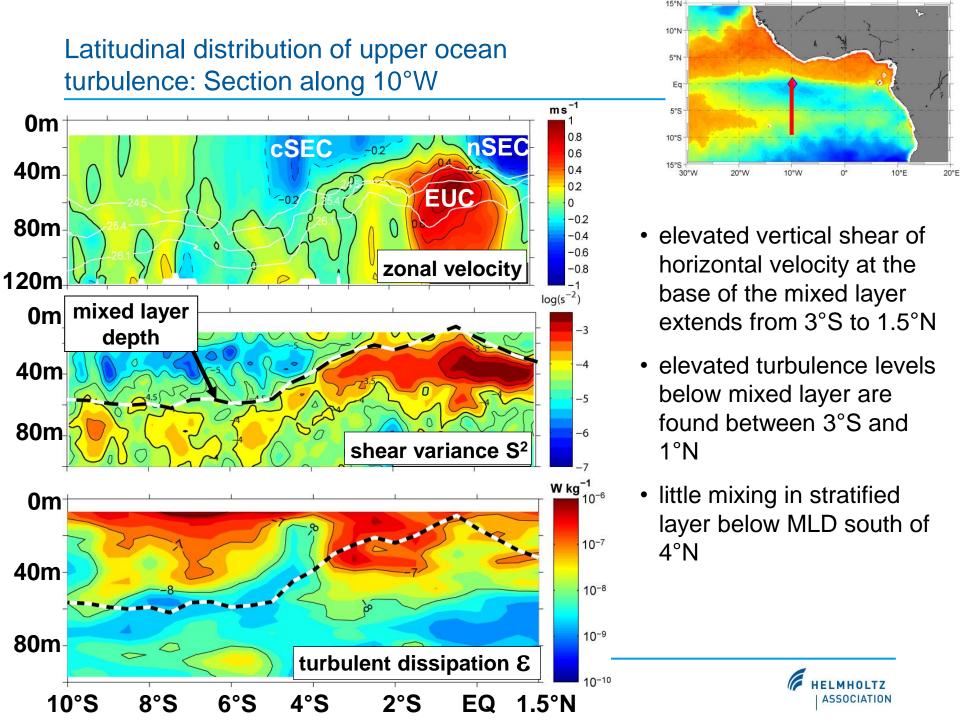




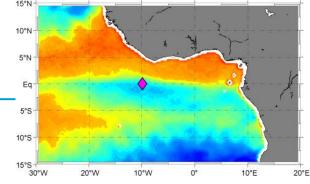


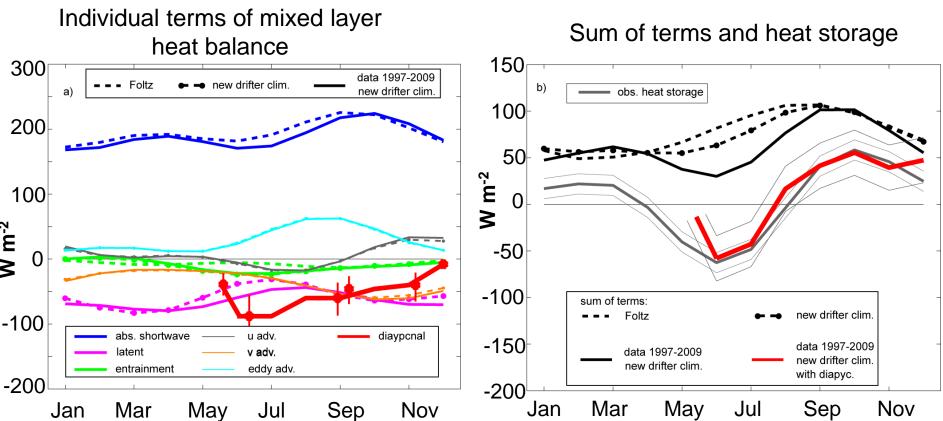
- microstructure probe (Rockland Scientific) attached to a Glider
- measures autonomously for up to 4 weeks
- profiles the water column to 1000m in about 45 minutes





Seasonal cycle of mixed layer heat budget at 0°N, 10°W

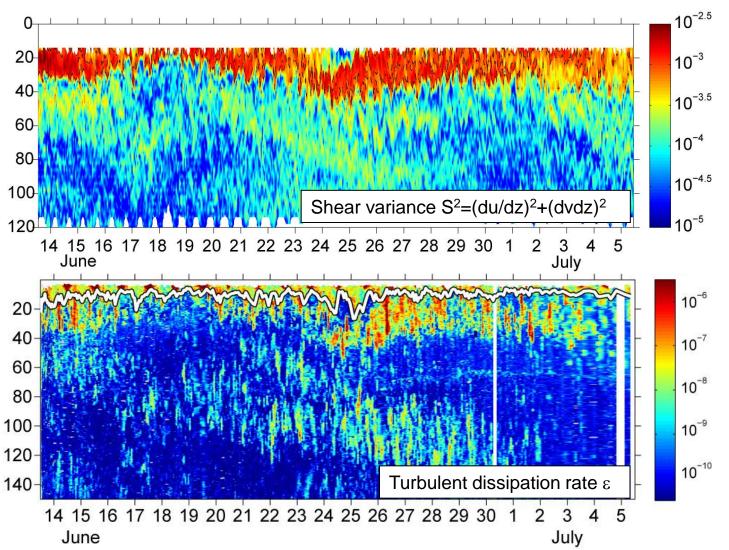


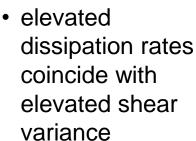


(Hummels et al., 2013)



Vertical shear of horizontal current and turbulent kinetic energy dissipation rates

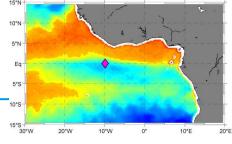


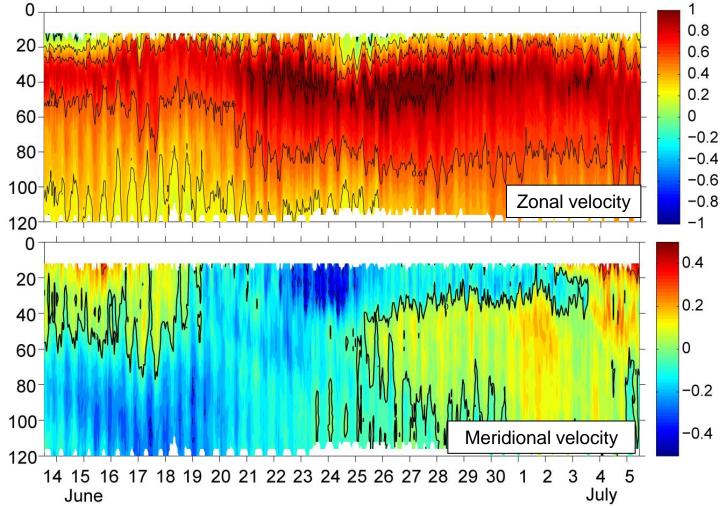


 bursts of elevated turbulence in the thermocline occur sporadically and last up to a few hours



Horizontal currents observed during the MircoRider/Glider mission

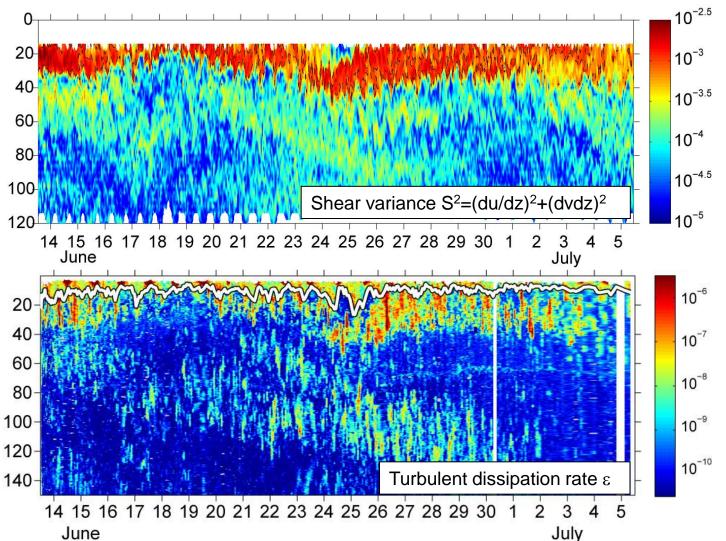


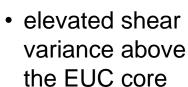


- Strong tidal currents in record with amplitude of ~8 cm s⁻¹
- core of the EUC located at 40m-60m depth



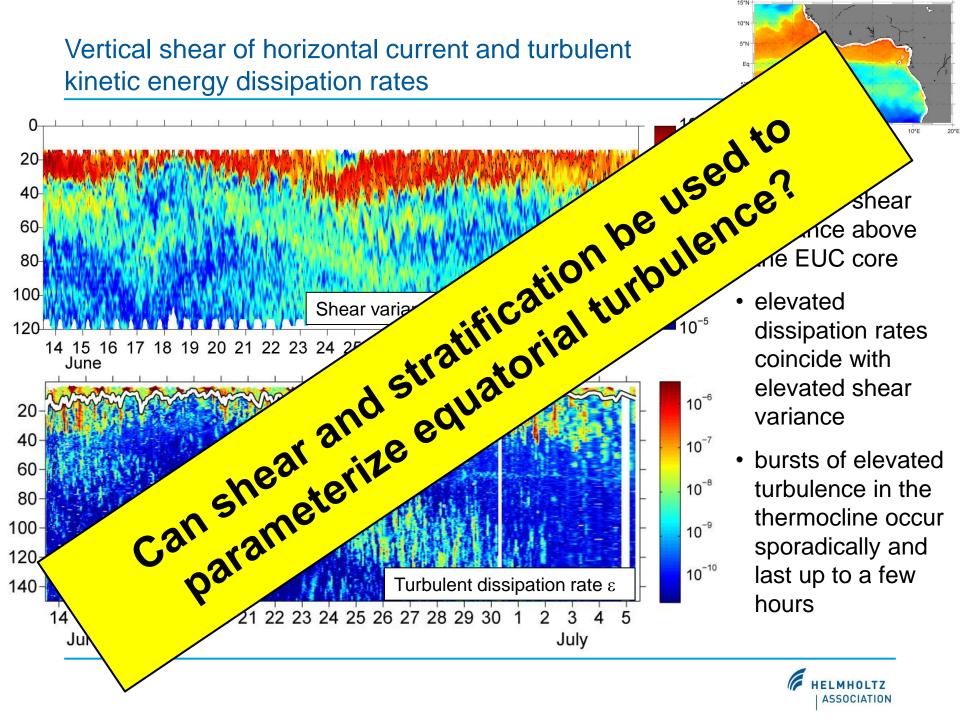
Vertical shear of horizontal current and turbulent kinetic energy dissipation rates





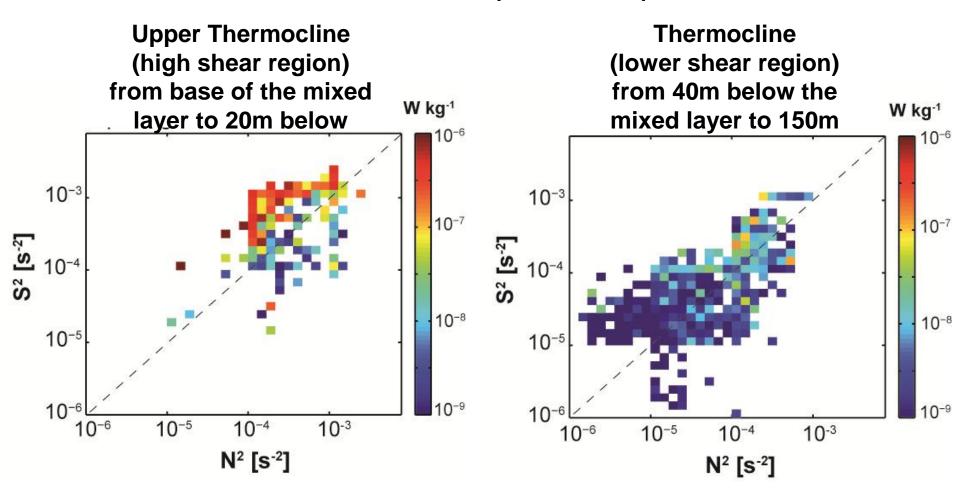
- elevated
 dissipation rates
 coincide with
 elevated shear
 variance
- bursts of elevated turbulence in the thermocline occur sporadically and last up to a few hours







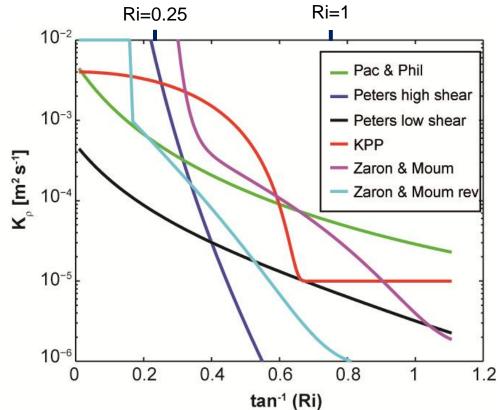
Observations (1.5°N – 2°S)





Ri-dependent mixing parameterizations





Pacanowski and Philander (1981)

$$v = \frac{50 \cdot 10^{-4} m^2 s^{-1}}{(1 + 5Ri)^2} + 10^{-4} m^2 s^{-1}$$

$$K_{\rho} = \frac{v}{(1+5Ri)} + 10^{-5} m^2 s^{-1}$$

Peters et al., (1988)

$$K_{\rho} = \frac{5 \cdot 10^{-4} m^2 s^{-1}}{(1 + 5Ri)^{2.5}} + 10^{-6} m^2 s^{-1}$$
 high shear

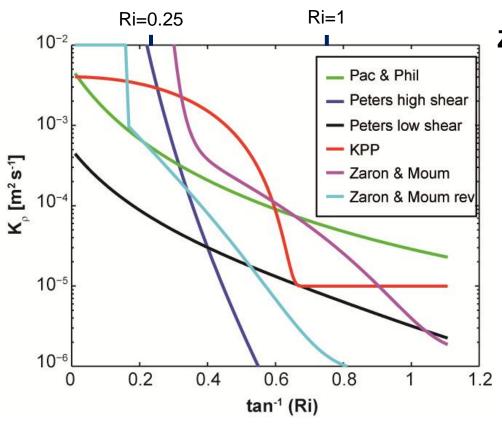
$$K_{\rho} = 1.1 \cdot 10^{-8} Ri^{-9.2}$$
 low shear

$$K_{\rho} = 50 \cdot 10^{-4} \, m^2 s^{-1} \left[1 - (Ri / 0.7)^2 \right]^3$$









Zaron & Moum (2009)

$$K_h^{alt} = \left| V \right|^2 / S \cdot a \left(\frac{Ri_1}{Ri - Ri_1} \right)^{\alpha} + be^{-\beta \cdot Ri} + c$$

$$K_h^{rev} = \left| V \right|^2 / S \cdot \Delta \phi_h e^{-\gamma (Ri - Ri_2)} + \phi_h^w$$

Uses additional parameter:

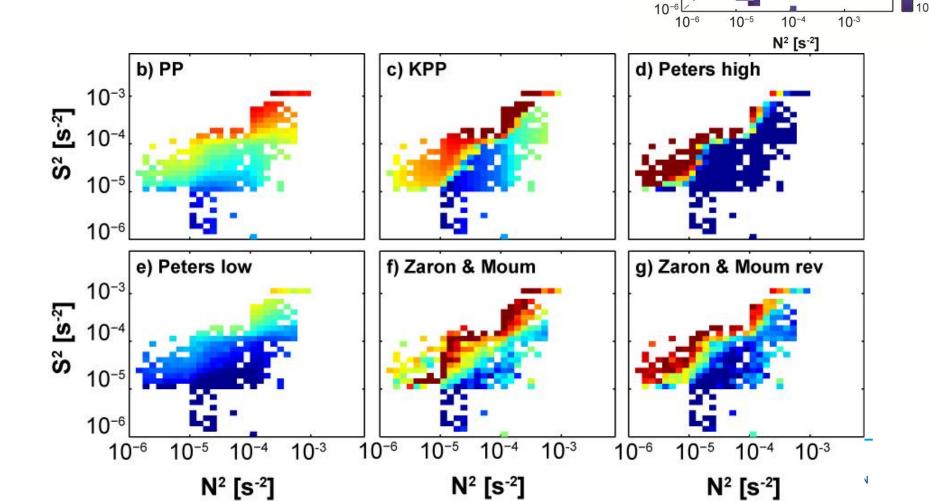
$$\left|V
ight|^2-$$
 large-scale kinetic energy $S-$ shear



equatorial data (2°S-1.5°N) upper thermocline W kg⁻¹ depth range: MLD-MLD+20 Average TKE dissipation rates in N² and S² bins 10^{-3} 10-7 5 10⁻⁴ 10-8 **Upper Thermocline (MLD to MLD+20m)** 10-5 10 10-5 10-6 10-4 10-3 N² [s⁻²] b) PP c) KPP d) Peters high 10-3 10-4 10^{-5} 10^{-6} e) Peters low f) Zaron & Moum g) Zaron & Moum rev 10-3 10-4 10-5 10^{-6} 10-3 10^{-5} 10-6 10-5 10^{-6} 10^{-4} 10⁻³ 10-6 10-5 10^{-4} 10-4 10^{-3} $N^2 [s^{-2}]$ N² [s⁻²] N² [s⁻²]

Average TKE dissipation rates in N² and S² bins

Thermocline (MLD+40m to 150m)



equatorial data (2°S-1.5°N)

depth range: MLD+40 -MLD+130

W kg⁻¹

10-6

10-7

10-8

thermocline

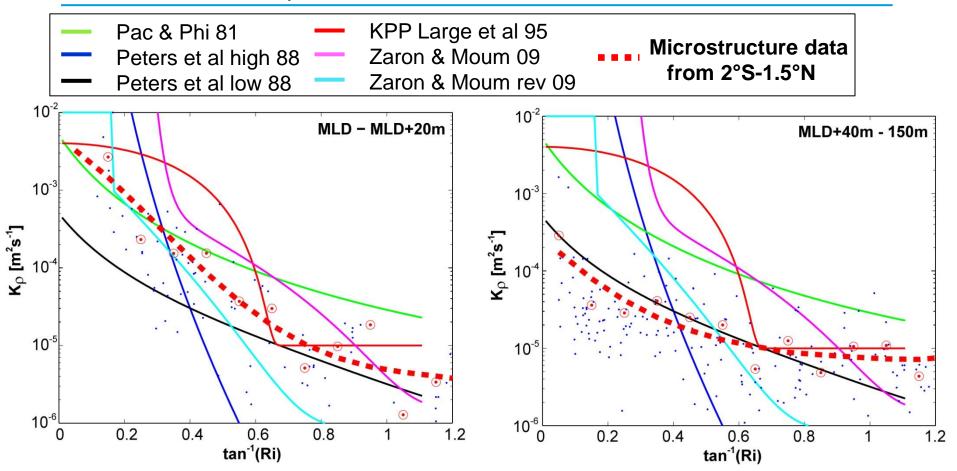
10-3

 10^{-5}

S² [s⁻²]



Evaluation of Ri-dependent Parameterizations







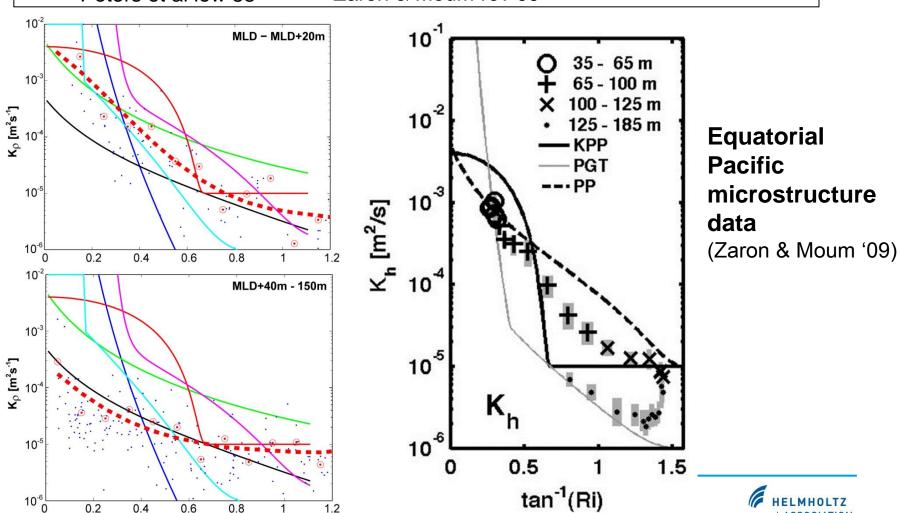
Evaluation of Ri-dependent Parameterizations

0.2

0.6

tan-1(Ri)

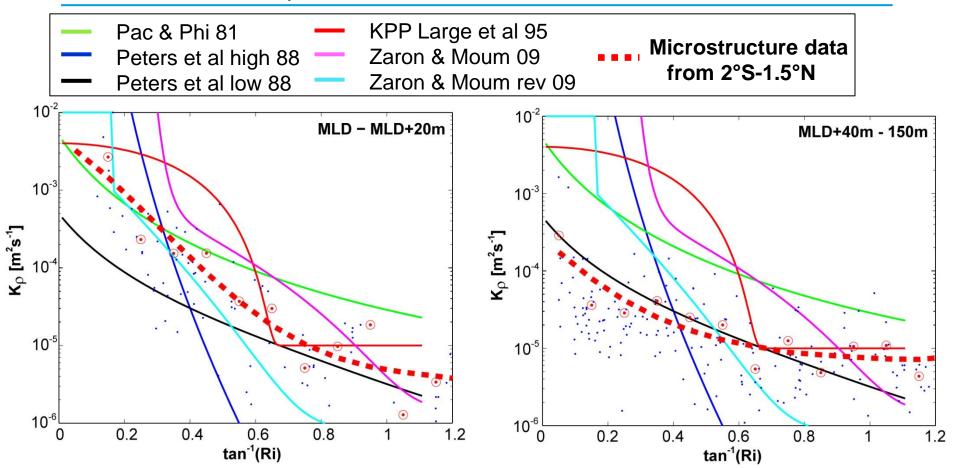




1.2



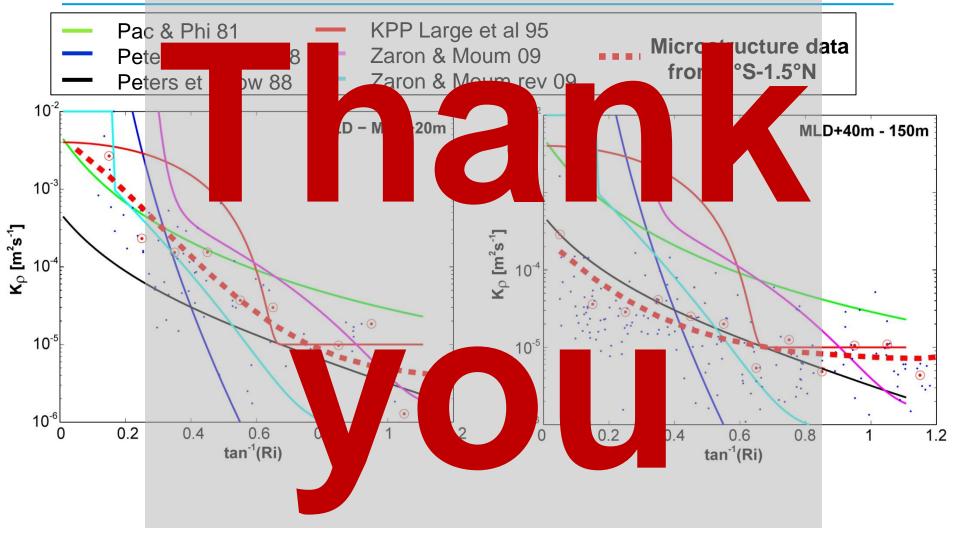
Evaluation of Ri-dependent Parameterizations





GEOMAR

Evaluation of Ri-dependent Parameterizations





Conclusions



existing Ri-dependent mixing parameterizations do not well reproduce





- Mixing Observatories at GEOMAR
- Parameterization of shear driven mixing in the tropical ocean
- Parameterization of internal wave driven mixing
- Conclusions

