

Supplementary Materials for
**Dominance of benthic fluxes in the oceanic beryllium budget and implications
for paleo-denudation records**

Kai Deng *et al.*

Corresponding author: Kai Deng, kai.deng@erdw.ethz.ch

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Supplementary Text

Estimation of continent ^{10}Be input to the oceans

The oceanic external ^{10}Be input includes atmospheric wet and dry deposition at the ocean surface, eolian dust, the riverine dissolved load, and a benthic flux via diagenetic release from continent-derived sediments. The major sink is scavenging onto marine particles, as for ^9Be . We consider atmospheric deposition as the sole dominant external source given the very small magnitude of other ^{10}Be sources. Here we aim to validate this assumption based on some simple calculations.

The atmospheric ^{10}Be deposition at the ocean surface ($5.77 \pm 1.41 \text{ mol/yr}$) is calculated by multiplying the global-average atmospheric ^{10}Be depositional flux (41) by the global ocean area. The effective riverine dissolved ^{10}Be input to the oceans is derived following three steps. 1) The atmospheric deposition to the global river basins ($1.56 \pm 0.38 \text{ mol/yr}$) is calculated as the product of the global-average atmospheric ^{10}Be depositional flux and the global drainage area (57), which sets the upper limit for continent-derived (dissolved and particulate) ^{10}Be input to the oceans. 2) Global riverine dissolved [^{10}Be] ($\sim 320 \pm 80 \text{ at/g}$) can be estimated by partitioning atmospheric ^{10}Be input into two components: the dissolved (river water) and particulate (river sediment) forms. Specifically, this partitioning is based on global water and sediment discharge data (57) and a Be partition coefficient ($1.5 \times 10^5 \text{ L/kg}$) (58) estimated from global discharge-weighted water pH (7.1) (59). 3) Hence, the effective riverine ^{10}Be input to the oceans ($0.009 \pm 0.002 \text{ mol/yr}$) is the product of the global riverine dissolved [^{10}Be], the global water discharge and the fraction of ^{10}Be escaping the estuarine trap (44%; same as ^9Be ; (14)). The dust ^{10}Be input ($0.15 \pm 0.05 \text{ mol/yr}$) is calculated as the product of the global dust flux (38) and dust [^{10}Be] represented by that in loess from the Chinese Loess Plateau (60), assuming that 100% of dust ^{10}Be is soluble, to yield an upper-limit estimate.

To estimate the diagenetic release of ^{10}Be from continent-derived terrestrial particles that mainly occurs on continental shelves, we first derive a global-average riverine sediment ($^{10}\text{Be}/^9\text{Be}$)_{reac} ($1.2 \pm 0.6 \times 10^{-9} \text{ at/at}$) based on its relationship with denudation rates (61) and the global average denudation rate of $233 \text{ t/km}^2/\text{yr}$ (57). Next, we estimate the benthic flux of ^{10}Be on continental shelves ($0.05 \pm 0.04 \text{ mol/yr}$) as the product of the ($^{10}\text{Be}/^9\text{Be}$)_{reac} above and the benthic ^9Be flux on continental shelves (Table S5). This might be an upper-limit estimate for the benthic ^{10}Be flux from terrestrial particles, because a small fraction of the shelf benthic Be flux might be sourced from the marine authigenic fraction rather than the terrestrial authigenic fraction (49).

To conclude, the effective continental ^{10}Be input to the oceans is only $\sim 0.21 \text{ mol/yr}$, and thus the atmospheric deposition accounts for 96.5% of the total external ^{10}Be input to the oceans. Considering that the relative uncertainty of atmospheric ^{10}Be deposition is 20-30%, adding or excluding other sources only slightly changes the estimate of the total external ^{10}Be input. Hence, we consider atmospheric deposition as the sole oceanic ^{10}Be source for simplification in the main text.

Table S1. Basic characteristics of sampling stations

Station	Continental margin	Longitude	Latitude	Water Depth	Bottom water O ₂ ¹
		°E	°N	m	μmol/L
C6-1	East China Sea	122.0	31.1	6	304
C10	East China Sea	122.5	31.0	12	306
C13	East China Sea	122.9	30.8	33	115
B14	East China Sea	122.9	30.1	46	203
AL543-8	Baltic Sea	10.2	54.5	23	<0.1
AL543-12	Baltic Sea	10.1	54.6	24	36
Soledad	Mexico margin	-112.7	25.2	542	<0.1
Catalina	California margin	-118.6	33.3	1300	19
San Clemente	California margin	-118.1	32.6	2053	52
Patton Escarpment	California margin	-120.6	32.4	3707	132

Note: 1) Bottom water O₂ data on the East China Sea shelf and the Mexico-California margins are from (50) and (26), respectively.

Table S2. Pore-water Be, Mn and Fe concentrations^{1,2,3,4}. Stations are shown in Table S1.

Station	Basin	Core depth cm	Be pM	Mn μM	Fe μM
C6-1	East China Sea	0	76.9	0.0	<D.L.
C6-1	East China Sea	1	130.2	4.9	<D.L.
C6-1	East China Sea	2	154.0	40.0	<D.L.
C6-1	East China Sea	4	165.4	47.6	<D.L.
C6-1	East China Sea	7	256.4	77.7	<D.L.
C6-1	East China Sea	9	255.8	99.5	<D.L.
C6-1	East China Sea	11	279.6	92.6	0.0
C6-1	East China Sea	17	411.4	110.3	0.0
C6-1	East China Sea	21	549.0	112.9	95.2
C10	East China Sea	0	46.3	<D.L.	0.0
C10	East China Sea	1	110.1	83.0	<D.L.
C10	East China Sea	3	166.5	75.3	<D.L.
C10	East China Sea	4	160.8	110.6	<D.L.
C10	East China Sea	6	193.7	162.7	0.1
C10	East China Sea	7	296.2	131.0	25.4
C10	East China Sea	9	343.8	73.0	211.1
C10	East China Sea	12	407.7	92.5	299.9
C10	East China Sea	16	401.2	89.8	337.6
C10	East China Sea	25	402.9	51.4	99.3
C13	East China Sea	0	45.8	0.0	0.1
C13	East China Sea	1	142.7	118.7	0.2
C13	East China Sea	2	278.7	162.3	94.7
C13	East China Sea	3	264.0	167.5	106.3
C13	East China Sea	4	268.6	92.1	105.2
C13	East China Sea	7	253.4	28.2	43.8
C13	East China Sea	9	225.0	26.0	60.1
C13	East China Sea	11	228.8	19.2	28.7
C13	East China Sea	17	222.4	21.9	22.8
C13	East China Sea	21	194.8	18.4	10.3
B14	East China Sea	0	54.8	0.0	<D.L.
B14	East China Sea	1	129.0	9.3	0.1
B14	East China Sea	2	282.2	20.0	30.5
B14	East China Sea	5	136.8	34.1	21.8
B14	East China Sea	6	222.3	46.9	55.2
B14	East China Sea	7	152.5	34.9	10.7
B14	East China Sea	9	156.2	27.6	3.7
B14	East China Sea	11	153.4	23.2	3.6
B14	East China Sea	15	145.1	20.3	5.9
B14	East China Sea	19	212.7	21.0	26.2
AL543-8	Baltic Sea	0	89.9	14.9	2.1
AL543-8	Baltic Sea	0.5	162.7	30.4	0.6

AL543-8	Baltic Sea	1.5	259.8	43.8	0.5
AL543-8	Baltic Sea	2.5	485.9	61.3	0.8
AL543-8	Baltic Sea	3.5	608.0	70.0	0.3
AL543-8	Baltic Sea	4.5	691.3	66.4	1.0
AL543-8	Baltic Sea	5.5	613.4	58.7	0.3
AL543-8	Baltic Sea	7	725.8	52.2	0.2
AL543-8	Baltic Sea	9	683.0	46.5	0.3
AL543-8	Baltic Sea	11	933.5	42.5	0.6
AL543-8	Baltic Sea	13	652.6	37.9	0.3
AL543-8	Baltic Sea	16	667.4	29.7	0.6
AL543-8	Baltic Sea	20	775.4	23.3	0.7
AL543-12	Baltic Sea	0	57.5	0.6	2.0
AL543-12	Baltic Sea	0.5	474.0	43.2	60.8
AL543-12	Baltic Sea	1.5	660.4	48.2	16.6
AL543-12	Baltic Sea	2.5	852.7	53.0	1.5
AL543-12	Baltic Sea	3.5	774.7	52.1	0.6
AL543-12	Baltic Sea	4.5	766.2	61.4	0.3
AL543-12	Baltic Sea	5.5	960.3	65.4	0.7
AL543-12	Baltic Sea	7	630.5	73.9	0.4
AL543-12	Baltic Sea	9	651.8	66.0	0.3
AL543-12	Baltic Sea	11		50.0	0.3
AL543-12	Baltic Sea	13	863.1	43.7	0.5
AL543-12	Baltic Sea	16	884.0	36.1	0.3
AL543-12	Baltic Sea	20	969.4	42.6	0.5
Soledad	Mexico margin	0.5	23.0	0.1	63.2
Soledad	Mexico margin	1.5	30.7	0.2	45.7
Soledad	Mexico margin	2.5	57.4	0.1	38.4
Soledad	Mexico margin	3.5	44.9	0.1	9.1
Soledad	Mexico margin	4.5	44.7	0.1	3.3
Soledad	Mexico margin	5.5	61.4	0.1	2.6
Soledad	Mexico margin	6.5	47.5	0.1	1.1
Soledad	Mexico margin	7.5	51.9	0.0	0.6
Soledad	Mexico margin	8.5	48.4	0.0	0.4
Soledad	Mexico margin	9.5		0.0	0.4
Catalina	California margin	0.5	16.8	3.3	0.0
Catalina	California margin	1.5	85.6	13.8	11.0
Catalina	California margin	2.5	110.9	6.1	19.0
Catalina	California margin	3.5	63.9	4.4	23.6
Catalina	California margin	4.5	87.8	3.5	24.9
Catalina	California margin	5.5	132.7	3.3	28.6
Catalina	California margin	6.5	109.4	3.4	30.4
Catalina	California margin	7.5	47.6	3.4	24.5
Catalina	California margin	8.5	83.5	3.0	23.0
Catalina	California margin	9.5	66.8	2.9	23.9

San Clemente	California margin	0.5	50.3	0.3	0.0
San Clemente	California margin	1.5	41.9	3.0	0.1
San Clemente	California margin	2.5	48.0	30.6	0.2
San Clemente	California margin	3.5	37.4	44.9	0.2
San Clemente	California margin	4.5	52.1	56.8	0.2
San Clemente	California margin	5.5	49.5	66.8	1.1
San Clemente	California margin	6.5	46.2	75.5	4.8
San Clemente	California margin	7.5	47.9	81.8	7.9
San Clemente	California margin	8.5	42.1	88.2	8.5
San Clemente	California margin	9.5	87.2	94.1	7.8
Patton Escarpment	California margin	0.5	52.2	0.0	<D.L.
Patton Escarpment	California margin	1.5	16.3	0.0	<D.L.
Patton Escarpment	California margin	2.5	47.5	0.0	<D.L.
Patton Escarpment	California margin	3.5	96.6	0.1	<D.L.
Patton Escarpment	California margin	4.5		1.0	0.1
Patton Escarpment	California margin	5.5	36.5	7.3	0.1
Patton Escarpment	California margin	6.5	34.9	11.2	0.1
Patton Escarpment	California margin	7.5	46.6	13.2	0.1
Patton Escarpment	California margin	8.5	41.1	13.2	0.1
Patton Escarpment	California margin	9.5		15.2	0.1

Note: 1) A relative uncertainty of 10% is applied for all element measurements.

2) <D.L.: below detection limit.

3) All element concentrations were measured at ETH Zürich using Element XR, except Mn data in the Baltic Sea that were measured at GEOMAR (Germany) using ICP-OES.

4) Fe-Mn concentration data from the East China Sea was previously reported in (50).

Table S3. Diffusive sedimentary Be fluxes from this study and the literature

Station	Basin	Lat. °N	Long. °E	Water depth m	T ¹ °C	Be gradient ² pmol/L/cm	Uncertainty pmol/L/cm	D _{Be} ^{sed3} cm ² /s	Diffusive Be flux ³ pmol/cm ² /yr	Uncertainty ⁴ pmol/cm ² /yr	Reference
C6-1	East China Sea	31.1	122.0	6	25	53.3	15.1	5.01E-06	7.6	2.2	This study
C10	East China Sea	31.0	122.5	12	25	63.7	11.9	5.01E-06	9.1	1.7	This study
C13	East China Sea	30.8	122.9	33	25	116.5	14.1	5.01E-06	16.6	2.0	This study
B14	East China Sea	30.1	122.9	46	25	113.7	14.4	5.01E-06	16.2	2.0	This study
AL543-8	Baltic Sea	54.5	10.2	23	14	158.4	19.8	3.69E-06	16.6	2.1	This study
AL543-12	Baltic Sea	54.6	10.1	24	13	833.1	95.5	3.62E-06	85.5	9.8	This study
Soledad	Mexico margin	25.2	-112.7	542	4	10.8	2.6	2.59E-06	0.8	0.2	This study
Catalina	California margin	33.3	-118.6	1300	4	36.9	6.1	2.59E-06	2.7	0.4	This study
San Clemente	California margin	32.6	-118.1	2053	4	39.9	11.7	2.59E-06	2.9	0.9	This study
Patton Escarpment	California margin	32.4	-120.6	3707	4	43.8	12.1	2.59E-06	3.2	0.9	This study
Site BC-10	California margin	32.5	-120.0	3800	4	46.3	8.5	2.59E-06	3.4	0.6	(19)
MANOP site S	North Pacific	11.0	-140.0	4910	4	19.7	3.0	2.59E-06	1.4	0.2	(19)

Note: 1) We adopted a bottom water temperature of 25 °C for the shallow East China Sea shelf based on long-term observation in summer (62). For slope-deep sea sites, we assume a uniform T of 4 °C.

2) This reflects the maximum gradient of pore-water [Be] to the sediment surface (bottom water). No bottom water Be concentrations were measured at the Mexico-California margin; for these sites, we thus used the bottom water [Be] in the Patton Escarpment (Site BC-10) from (19). The choice of this concentration (30 pM) is reasonable because deep sea [Be] is relatively stable, e.g. 28 pM in the North Pacific (55).

3) D_{Be}^{sed} : Be diffusion coefficients in sediment adjusted for bottom water temperature (36). Diffusive Be flux here is calculated using Eq. (1).

4) The uncertainty of the diffusive flux is propagated from that of the Be gradient.

Table S4. Comparison between globally-representative and our studied range of key environmental variables.

Variable ^{1,2,3}	Unit	0-0.2 km		0.2-4 km		>4 km		Reference	
		Globe	Study region ⁴	Globe	Study region ⁴	Globe	Study region ⁴	Globe	Study region ⁴
Water depth	km	<i>0.065</i> (0-0.2)	(0.012-0.046)	2.8 (0.2-4)	(1.3-3.7)	4.9 (4->10)	4.9	(37)	\
Sedimentation rate ⁵	mg/cm ² /yr	<i>496</i>	<i>380</i> (200-5000)	1.73 ± 1.73	(3-15)	1.08 ± 0.97	(0.06-<0.6)	0-0.2 km: (37, 63); >0.2 km: (64)	(65); (26); (66)
Org. C rain rate	mmol C/m ² /d	<i>13.4</i>	<i>15.2</i> (4.4-42.7)	<i>1.4</i>	(0.5-2.5)	<i>0.16</i>	0.13 ± 0.11	(67)	(68); (26); (69)
Org. C remineralization rate	mmol C/m ² /d	<i>9.4</i>	8.4 (3.6-17.6)	<i>1.2</i>	(0.4-1.3)	<i>0.15</i>	0.13 ± 0.11 ⁶	(67)	(68); (26); (69)

Note: 1) The value of each environmental variable is displayed as its average value ± standard deviation (minimum-maximum). As the data were reported in different forms in literature, certain statistical parameters are not applicable and thus not shown in some cases.

2) These variables on sedimentary and diagenetic environments are chosen because they might modulate benthic sedimentary Be fluxes.

3) Parameter values in italic indicate area-weighted average values.

4) The range of values represents that in the study region (East China Sea shelf, Californian margin or deep North Pacific) chosen for spatial extrapolation (Table S5). References for each study region are listed as: 0-0.2 km; 0.2-4 km; >4 km.

5) The area-weighted sedimentation rate on global shelves is calculated using the empirical relationship between water depth and sedimentation rate (in cm/yr) (63), a bulk dry density (1.2 g/cm³) identical to that in the study region (20), and the bathymetry (37). The average sedimentation rate and its uncertainty at depths > 0.2 km over global scale are sourced from globally compiled ²³⁰Th normalized sedimentary mass flux data (64). The sedimentation rate of the study region at >4 km (~1-<10 mm/kyr; (66)) is converted to mg/cm²/yr using a depth-integrated bulk dry density of ~0.6 g/cm³ (70).

6) The organic carbon remineralization rate (remineralization rate = rain rate - burial rate) in the study region at >4 km is set to be equal to the measured rain rate, as the organic carbon burial rate is estimated to be <0.001 mmol C/m²/d (69).

Table S5. First-order estimate of oceanic Be budget

Be source or sink	[Be] pmol/kg	Material flux kg/yr	Contributing area ¹⁰ km ²	Be input or output 10 ⁷ mol/yr	Proportion %	Reference ¹
River dissolved load ^{1,2,3}	514 ± 104	3.6E+16	9.8E+07	1.8 ± 0.4	26%	(57), Table S6, (14)
Eolian dust ^{1,2,4}	2.0E+07	(4.4 ± 0.3)E+11	3.6E+08	0.9 ± 0.06	12%	(38), (39), (13)
Benthic flux (shelf) ^{5,8}	140.0 ± 107.3 pmol/cm ² /yr		2.7E+07	3.8 ± 2.9	52%	This study
Benthic flux (0.2-4 km depths) ^{6,8}	2.9 ± 0.3 pmol/cm ² /yr		1.4E+08	0.4 ± 0.02	6%	This study
Benthic flux (>4 km depths) ^{7,8}	1.4 ± 0.2 pmol/cm ² /yr		1.9E+08	0.3 ± 0.03	4%	(19)
Be source (total) ¹¹				7.2 ± 2.9	100%	
Be sink (sediment burial) ^{9,11}				6.9 ± 1.7	\	(41), (10)

Note: 1) The references for river or dust source are listed as: material flux, [Be], delivery efficiency to the oceans.

2) [Be] here is calculated by accounting for the corresponding Be delivery efficiency (estuarine transmission of 44%; dust solubility of 9%).

3) The relative uncertainty of riverine [Be] is estimated as the $(\text{Be})_{\text{riv-diss}}$ flux-weighted relative standard deviation of [Be] from the compiled dataset in Table S6.

4) The material flux for dust is calculated from the average value and the standard deviation of three model results presented in (38) and compiled by (71).

5) The shelf benthic flux (pmol/cm²/yr) is represented by the advection-dominated flux (using Eq. 3) on the East China Sea shelf (details in “Materials and Methods”). The flux uncertainty is derived from the standard deviation of Be gradients at these stations and the uncertainty on water exchange rate (Fig. S1).

6) The benthic flux (0.2-4 km) is estimated by the diffusion-dominated flux in the California margin. To obtain this value, we use the average diffusive flux from three stations at 1300-3707 m with non-anoxic conditions. The flux uncertainty is derived from the standard deviation of the fluxes at these stations.

7) The benthic flux (>4 km) and its uncertainty are estimated using the diffusion-dominated flux at a station in the deep North Pacific (4910 m; (19)).

- 8) The global benthic ^9Be input is 4.5×10^7 mol/yr, accounting for only $\sim 0.9\%$ of the total riverine sediment-derived ^9Be input (5.3×10^9 mol/yr) (57, 61).
- 9) The sediment burial output of ^9Be is calculated by dividing the atmospheric ^{10}Be input to the oceans, that is, the product of global grid-average ^{10}Be depositional flux in atoms/ km^2/yr (41) and ocean area in km^2 , by the global ocean area-weighted deep-ocean $^{10}\text{Be}/^9\text{Be}$ ratio (10) following the approach of (40). The uncertainty of the ^9Be burial flux reflects propagated uncertainties of both parameters.
- 10) The contributing area refers to the global drainage area for river dissolved load (57) or ocean area for other sources and the sink (37).
- 11) The mismatch between oceanic fluxes of the sources and the sink of Be is $< 10\%$ and falls within the uncertainty.

Table S6. Compilation of global riverine dissolved Be concentration data

River Basin	Reference	Water discharge m ³ /yr	[Be] _{diss} pmol/L	[Be] _{diss} SD ¹ pmol/L
Amazon	(13, 72, 73)	6.46E+12	1049	219
Orinoco	(73)	1.13E+12	923	341
Ganges	(13)	4.92E+11	42	
Yangtze	(72)	9.27E+11	128	
Congo	(72)	1.23E+12	3824	625
Pearl	(74, 75)	1.67E+11	233	9
Mackenzie	(52)	2.84E+11	316	
Ob	(52)	4.26E+11	1598	
Yenisey	(52)	6.24E+11	317	
Lena	(52)	6.02E+11	1176	
Hudson	(75)	1.50E+10	2000	
Columbia	(75)	1.80E+10	353	
Fraser	(75)	8.60E+10	1055	49
Pitt	(75)	1.70E+09	215	
Santa Ana	(75)	7.40E+07	187	
Cosumnes	(75)	4.50E+08	2340	
Mokelumne	(75)	9.10E+08	1980	
Stanislas	(75)	7.30E+08	87	
Deer Creek	(75)	3.50E+07	686	
Merced	(75)	5.60E+08	261	
Klamath	(75)	3.70E+09	233	
Teviot	(76) ²	6.19E+08	2219	
Tweed	(76) ²	2.51E+09	2219	
L. Swale	(76) ²	6.46E+08	4439	
Ure	(76) ²	6.47E+08	2219	
Nidd	(76) ²	4.65E+08	3329	
Ouse	(76) ²	1.52E+09	2219	
Wharf	(76) ²	5.46E+08	2219	
Derwent	(76) ²	5.47E+08	2219	
Aire	(76) ²	1.15E+09	2219	
Calder	(76) ²	5.62E+08	3329	
Don	(76) ²	1.77E+08	5548	
Trent	(76) ²	2.86E+09	2219	
Upper River Severn	(77)	1.67E+07	8877	
Vörå creek	(78)	1.12E+08	269363	11848
Awe	(14)	1.80E+09	494	
Etive	(14)	4.00E+08	610	
Krka	(14)	1.50E+09	92	
Zrmanja	(14)	1.20E+09	54	

Continued below

Summary		
Parameter	Unit	Value
Sum of discharge	m ³ /yr	1.25E+13
Discharge-weighted [Be] _{diss}	pmol/L	1167
Global water discharge	m ³ /yr	3.60E+13
Global (Be) _{riv-diss} flux ³	mol/yr	4.20E+07
(Be) _{riv-diss} flux (post-estuary) ⁴	mol/yr	1.85E+07
Flux-weighted uncertainty ⁵	mol/yr	3.76E+06

Note: 1) If [Be]_{diss} data of multiple samples (spatial/temporal replicates) are available, we calculate the average value and the standard deviation (SD) in Columns "[Be]_{diss}" and "[Be]_{diss} SD", respectively.

2) Data in (76) was originally reported in units of µg/L and thus its precision is likely close to ~0.01 µg/L (~1 nM).

3) Global riverine dissolved Be flux is calculated using the global water discharge (57) and the discharge-weighted [Be]_{diss} of our compiled dataset.

4) Riverine dissolved Be flux after estuarine scavenging is calculated using a global-average scavenging efficiency of 56% (14).

5) The relative uncertainty of riverine dissolved Be flux (post-estuary) is calculated using the existing data on relative standard deviation of [Be]_{diss} (Column "[Be]_{diss} SD") weighted by their riverine Be flux.

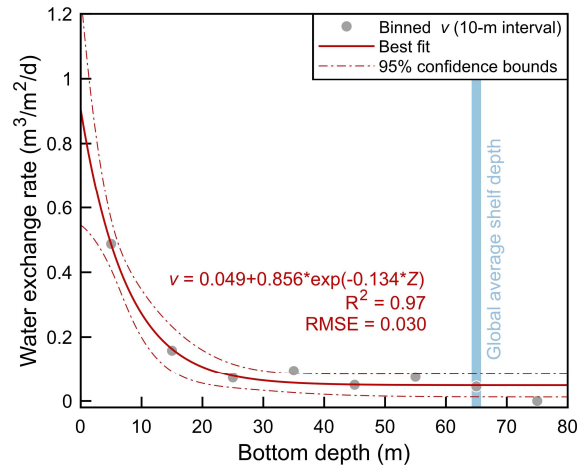


Figure S1 Relationship between bottom water depth (Z) and water exchange rate (v) on continental shelves. Water exchange rate data in the literature were derived from $^{224}\text{Ra}/^{228}\text{Th}$ disequilibria in China coastal seas (33). Data were binned at 10 m depth interval and fitted using an exponential equation following (33). Note that the literature data (33) were re-processed here in order to derive an estimate of uncertainty (red dashed lines) on v , which was not available in the original paper. At global average shelf depth, the water exchange rate is $0.049 \pm 0.036 \text{ m}^3/\text{m}^2/\text{d}$. The uncertainty on v represents the overall deviation from the best-fit equation and takes the regional heterogeneity of v into account.

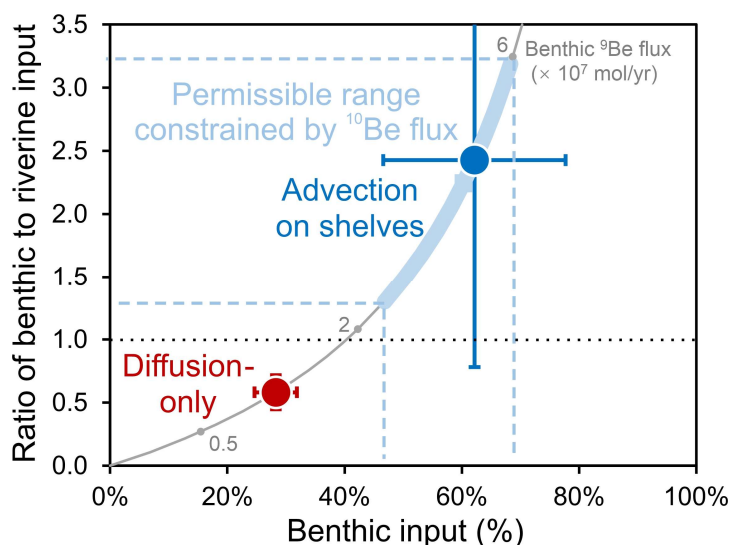


Figure S2 Sensitivity of benthic ${}^9\text{Be}$ input (% of the total input; X axis) and ratio of benthic to riverine ${}^9\text{Be}$ input (Y axis) to different scenarios on benthic flux calculation. Error bars include all the uncertainties propagated from every ${}^9\text{Be}$ source or sink. The permissible range constrained by the ${}^{10}\text{Be}$ flux (light blue band) indicates the range of benthic ${}^9\text{Be}$ fluxes ($4.2 \pm 1.7 \times 10^7$ mol/yr) calculated from input-output balance of ${}^9\text{Be}$, that is, the difference between the total ${}^9\text{Be}$ output (constrained by the global atmospheric ${}^{10}\text{Be}$ input and the global average deep-sea ${}^{10}\text{Be}/{}^9\text{Be}$; Table S5) and non-benthic ${}^9\text{Be}$ sources (river and dust; Table S5). This permissible range of benthic ${}^9\text{Be}$ flux is then converted to a range of paired X-Y values by comparing to fluxes of other non-benthic sources. The benthic ${}^9\text{Be}$ flux calculated by only diffusion processes (i.e. advection on shelves not included; red circle) accounts for 28% of the total input and corresponds to ~60% of the riverine dissolved input. However, this estimate does not fall within the permissible range of the benthic source. In comparison, the benthic ${}^9\text{Be}$ flux that takes advection processes on shelves into account (dark blue circle) agrees well with the permissible range. Note that even when considering the large flux uncertainty related to estimates on water advection, the total benthic flux is close to, or higher than, the riverine dissolved input (minimum ratio: 0.8; average ratio: 2.4), and accounts for $62\% \pm 16\%$ of the total ${}^9\text{Be}$ input.

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